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Coseismic dehydration of serpentinite: Evidence from high-velocity friction experiments

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25 **ABSTRACT**

26 High-velocity friction (HVF) experiments on serpentinite under conditions equivalent
27 to large amounts of earthquake slip produced large volumes of water vapor derived from
28 the frictional-heating-induced dehydration of serpentinite. Fourier Transform Infrared
29 (FTIR) and thermogravimetry (TG)-Differential Scanning Calorimetry (DSC)
30 microspectroscopic analyses show that the water in the slip zone and its bounding zones
31 was released due to dehydration during the HVF experiments. Our experimental results
32 demonstrate that (i) the run product consists of ~10 wt.% molten materials with
33 abundant vesicles and ~90 wt.% fine-grained clasts; (ii) both serpentine and olivine
34 minerals were melted during high-velocity slip (1.1–1.3 m/s); (iii) rapid serpentine
35 dehydration occurred in a zone of up to ~3 mm wide, including the slip zone and its
36 bounding zones, accompanying frictional melting during high-velocity slip; and (iv) no
37 distinct dehydration or frictional melting occurred during low-velocity slip (<0.4 m/s).
38 These findings show that dehydration reactions of serpentine can be caused by frictional
39 heating that accompanies frictional melting in the slip zone, and by the rapid diffusion
40 of frictional heat from the main slip zone to a wide zone bounded by the coseismic slip
41 plane during seismic faulting. The present results reveal that the thermal pressurization
42 caused by the dehydration of serpentine and frictional melting is a main mechanism that
43 may lead to the dynamic weakening of seismogenic faults, thereby facilitating seismic
44 slip during large earthquakes in subduction zones and along intracontinental faults that
45 contain abundant hydrous minerals.

46

47 **Keywords:** dehydration, serpentinite, high-velocity slip, fault strength, dynamic

48 weakening of seismogenic fault, thermal pressurization

49

50 1. Introduction

51

52 Thermal pressurization has been proposed as a major mechanism of lowering the
53 dynamic strength of faults during earthquake slip [e.g., Sibson, 1973; Lachenbruch,
54 1980; Andrews, 2002; Wibberley and Shimamoto, 2005]. A commonly invoked
55 mechanism of thermal pressurization in this regard is a reduction in effective normal
56 stress induced by the sudden increase in pore pressure that arises from the release of
57 fluid via dehydration reactions [e.g., Meade and Jeanloz, 1991; Tibi et al., 2002]. In
58 recent years, many experimental studies have suggested that thermal pressurization may
59 result from mineral decomposition, such as the dehydration of serpentine or kaolinite
60 [e.g., Hirose and Bystricky, 2007; Brantut et al., 2008; Viti and Hirose, 2010; Kohli et
61 al., 2011], the decarbonation of calcite or siderite [Han et al., 2007, 2010], and the
62 dehydration of gypsum [Brantut et al., 2011]. Observations of natural seismogenic fault
63 cores that contain fault gouges also support the supposition that dehydration reactions
64 occur during seismic faulting [e.g., Famin et al., 2008; Hamada et al., 2009, 2011].

65 Serpentine is closely associated with transform faults [e.g., Christensen, 1972;
66 Francis, 1981; Moore and Rymer, 2007] and with seismogenic faults within subduction
67 zones [Ulmer and Trommsdorff, 1995]; indeed, the presence of serpentine minerals is
68 considered to facilitate creep along the San Andreas Fault [Moore and Rymer, 2007],
69 thereby explaining its low fault strength [Wibberley, 2007]. In numerical modeling, heat
70 diffusion is considered an important coupling factor of thermal pressurization, which

occurs normal to the slip zone [Brantut et al., 2010]. Should serpentine dehydration occur due to frictional heating along seismogenic faults and should heat diffusion occur within a wide dehydration reaction zone bounded by the slip zone during earthquake slip, it is likely that a considerable volume of water would be released from serpentine-rich slip zones and their bounding zones. Thus, it is proposed that a rapid dehydration reaction involving serpentine would induce a sudden increase in fluid pressure that would simultaneously act to reduce the effective normal stress and markedly weaken the dynamic strength of seismogenic faults during seismic faulting.

To test this hypothesis, a low- to high-velocity rotary shear test apparatus [Lin, 2008] was used to conduct uniaxial high-velocity friction (HVF) experiments under conditions equivalent to large earthquake slip with various slip rates, using natural serpentinite samples from Nagano Prefecture, Japan. This paper describes the results of these experiments, which showed the instantaneous dehydration of serpentinite due to frictional heating along the simulated fault and heat diffusion within a wide zone bounded by the slip zone.

2. Experimental Procedure and Conditions

2.1. Test Equipment

All of the present experiments on frictional-heating-induced hydration of serpentine were performed under dry conditions using a rotary-shear low- to high-velocity frictional test machine installed at Shizuoka University, Japan (Fig. 1). This machine

contains two testing devices: a uniaxial rotary-shear low- to high-velocity friction device that lacks a hydraulic pressure vessel [as used by Lin and Shimamoto, 1998], and a hydraulic friction apparatus with a controlling device for high pore-water pressure, attached to the uniaxial rotary-shear low- to high-velocity friction device (Figs 1 and 2). This device has a wide range of slip rates, from <10 cm/yr to 10 m/s, yielding mechanical data such as torque, the frictional coefficient, amount of axial shortening, normal stress and shear stress, pore-water pressure, frictional melt temperature, and displacement. The test equipment can be used to apply normal stresses up to approximately 100 MPa, close to the typical uniaxial strength of crystalline rocks, provided that anvil-shaped specimens are used. In practical use, however, it is only possible to apply normal stresses of <15 MPa because of the severe thermal fracturing of unconfined specimens that occurs upon frictional heating.

2.2. Experimental Conditions

Serpentinite samples for the experiments, collected from Nagano Prefecture, central Japan, were free of visible cracks and consisted of $\sim 65\%$ antigorite and $\sim 35\%$ olivine, along with trace amounts of mafic minerals (Fig. 3).

Ten HVF experiments were performed at room temperature under dry conditions and with equivalent slip rates [Shimamoto and Tsutsumi, 1994; Lin and Shimamoto, 1998] of 1.10–1.31 m/s and normal stresses of 2–9 MPa (Table 1). In such experiments, the slip rate varies from the inner to outer circumferences of the sample, so we used the equivalent slip rate (V_{eq}), which is defined as V_{eq} multiplied by the area of sliding

equals the rate of frictional work. Assuming a constant frictional coefficient over the sliding surface, V_{eq} for a cylindrical specimen of diameter r is given by

$$V_{eq} = 4/3 \pi r R$$

where R is the revolution rate of the motor [Shimamoto and Tsutsumi, 1994].

Low-velocity (0.35–0.4 m/s) friction (LVF) experiments were conducted for comparison with the results of the HVF experiments (Table 2). A pair of cylindrical specimens (25 mm in diameter and 50 mm in length) was used in each experiment, with the circular interface between the two specimens being set as the slip plane (Fig. 2). The sliding surfaces were polished with 1500# SiC powders. Because large normal stresses could not be applied to the simulated fault due to the limitations of the experimental apparatus, the HVF experiments were run over a period of ~15–40 sec, corresponding to displacements of 20–50 m. These displacements are comparable with those recorded during large-magnitude earthquakes (e.g., maximum displacements of 15–30 m have been reported for the 2001 M_w 7.8 Kunlun earthquake [Lin et al., 2002], the 2004 M_w 9.0 Sumatra–Andaman earthquake [Subarya et al., 2006], and the 2011 M_w 9.0 Offshore Tohoku (Japan) earthquake [Yagi and Fukahata, 2011]). In contrast, the total displacements over a period of ~15–40 sec in the LVF experiments were calculated to be 5–15 m. The frictional energy expended per unit area in the experiments under a slip rate of ~1 m/s and a large displacement of up to 50 m is similar to that expended during a medium- to large-magnitude crustal earthquake [O’Hara et al., 2006], at least in terms of slip velocity and stress drop [Kanamori and Brodsky, 2004].

Serpentine dehydration phenomena were directly observed during the HVF experiments (1.10–1.31 m/s) through a transparent window cover on the experimental

apparatus (Fig. 4). The simulated fault began to produce water vapor and dust within 0.2 sec of the initiation of fault motion (Fig. 4a). Water vapor, accompanied by sparks and red melt material, was sprayed out from the slip zone, with some instantaneously cooling to water droplets (0.1–5.0 mm in diameter) on the transparent window cover (Fig. 4b–d). The ejection of water vapor from the simulated fault continued until the end of each experiment, with the droplets on the window cover forming a band of 3–7 mm in width (Fig. 4b–d). In contrast, such water vapor and water droplets were not observed during the LVF experiments.

3. Results

3.1. Microstructures of Run Products

The slip zones that formed during the experiments ranged in width from 0.1 to 2.0 mm, consisting of numerous fractured fragments of olivine within a very fine-grained matrix that was too fine to enable the identification of minerals (Figs 5 and 6), even by powder X-ray diffraction analysis. The olivine crystals abutting the slip zone were generally fractured, with numerous cracks infilled by molten run products containing fine-grained clasts, forming complex network veins branching from the main slip zone (Fig. 5a). These veins resemble natural pseudotachylite veins in their geometry and morphology [Lin, 2008]. In contrast, the antigorite crystals that abut the slip zone were mainly decomposed, without distinct cracks or fibrous textures as those observed in the host rock far from the slip zone (Figs 5 and 6a–c). These textural features indicate that

the serpentine crystals were decomposed in a narrow zone of <3 mm width, developed each side of the slip zone, within which the serpentine was subjected to dehydration reactions during the HVF experiments (see below for details).

Some molten run products were extruded from the simulated fault after 0.5 sec of the onset of slip, occurring as irregular branch-like forms of <2 mm in diameter and up to 3 cm in length, showing flow structure and with a vitreous luster similar to that of fresh lava (Fig. 7a). Molten material that includes numerous fine-grained clasts is generally heterogeneous in texture under the microscope, showing flow structure and containing numerous spherical to elliptical vesicles ranging in diameter from sub-micron size to 0.5 mm (Figs 6d and 7b). Such vesicular structures are considered to form via the extrusion of gas (water vapor) from a melt during frictional melting, as reported previously in both natural pseudotachylyte [e.g., Maddock et al., 1987; Lin, 1994, 2008; Magloughlin, 2011] and experimentally generated pseudotachylyte [Lin and Shimamoto, 1998]. These structural features observed in the present experiments indicate that molten material formed and water vapor was released from the slip zone during the HVF experiments (Fig. 4).

The vesicles and flow structures are observed in narrow zones of <3 mm wide bounded by antigorite crystals on each side of the slip zone (Figs 5a and 6a–c). The fibrous textures of antigorite crystals observed in the starting serpentine samples are not recognized in the dehydration zones that abut the slip zone (Figs 5b and 6a–c).

3.2. Powder X-ray Diffraction Analysis

Two types of run products were analyzed by powder X-ray diffraction method: one is the molten material remained in the slip zone (HVF R001), and another is the molten material of run product extruded from the slip zone during the HVF experiment (HVF R012). For comparison with the run products, the host serpentinite sample and olivine mineral sample were also analyzed. Powder X-ray diffraction analyses reveal that the diffraction peaks of antigorite are absent from the spectra of the run products, even though antigorite made up ~65% of the starting samples (Fig. 8). The integrated intensities of olivine peaks in the spectra for run materials (Fig. 8c, d) are slightly stronger than those for the serpentinite sample (Fig. 8a). The diffraction spectra of both types of run products contain a distinctive, broad band ranging from 2θ values of 12° to 42° , comparable with that for a standard sample containing 5–10 wt.% glass (Fig. 8d–f). These diffraction patterns are also comparable with the XRD spectra of standard samples analyzed for the quantitative calibration of glassy materials contained in the HVF experimentally generated pseudotachylyte, as reported by Lin and Shimamoto (1998). This finding indicates that up to 10 wt.% glass or amorphous material formed in the run products that were extruded and remained in the slip zone during the present experiments, as observed under optical and electron microscopes (Figs 5 and 6).

3.3. Chemical Compositions

The bulk chemical composition of serpentinite samples used in the frictional experiments was analyzed by X-ray fluorescence (XRF) (sample Host1 in Table 2). For comparison with the run products, the chemical composition of serpentine and olivine

minerals in samples of host rock (samples HS1-HS5 and HO1-2 in Table 2) was measured using an electron-probe microanalyzer (EPMA, JXA733). The chemical compositions of run products in the slip zone and in the dehydration zones (Fig. 5) were measured from a thin section oriented perpendicular to the slip zone, as shown in Fig. 3c (Tables 3–4).

The molten materials derived from the slip zone are generally heterogeneous in chemical composition (Fig. 9, Table 3), which differs from the bulk composition of the host rock, and from that of serpentine and olivine minerals contained in the host rock, respectively (Fig. 9, Table 2). The serpentine minerals in the dehydration zones (samples Dh1–Dh8 in Table 4) have a similar composition with that of serpentine minerals in the host rock with the sum of elements of 2.82–2.84 (Fig. 9, samples HS1–HS6, in Table 2). The variation diagrams of major oxides show a characteristic change in SiO₂, MgO, FeO, CaO and Al₂O₃ components in the run products (Fig. 9). Comparing with the dehydration zones, the molten materials in the slip zone have lower SiO₂ and Al₂O₃ components and higher FeO component similar to that of olivine minerals contained in the host rock (Fig. 9). The slip zone has a large variation range in CaO and FeO components which are generally higher than that of serpentine minerals contained in the host serpentinite. There is also a large variation range in MgO component of the molten materials. In contrast, there is a small variation range in MgO component in serpentine minerals contained in the host serpentinite and dehydration zone (Fig. 9). The variations of chemical composition in the molten materials may be caused by friction melting of serpentine and olivine minerals, in which molten materials were not completely mixed and contaminated.

The analytical totals of the molten materials in the slip zone and serpentine minerals that about the slip zone are up to >97.6–99.6 wt.% (Tables 3–4), higher than that of serpentine minerals in the host rock (samples HS1–HS6 in Table 2). These high total amounts can be interpreted as a result of water release from the slip zone and bounding zones due to the dehydration of serpentine minerals (see below for details).

3.4. Frictional Coefficient and Shortening Amount

The frictional coefficient μ showed an immediate increase to 0.56 at the onset of slip, within the first 0.1 sec of the experiments, but just as quickly dropped to a steady-state value of 0.23 prior to an elapsed time of 0.2 sec during the HVF experiments (Fig. 10a–b). This pronounced peak in friction accompanied the outspray of water vapor and axial shortening of the sample within the first 0.2 sec of each run. In contrast to the HVF experiments, during the LVF experiments the frictional coefficient μ increased to 0.57 within a 2.5-sec period after 1 sec of the onset of slip, and gradually dropped to a steady-state value of 0.28 prior to an elapsed time of ~20 sec, without distinct axial shortening in the first 60 sec (Fig. 10c–d). This result indicates that slip weakening occurred over a shorter period during the HVF experiments compared with the LVF experiments.

During the HVF experiment, axial shortening occurred in the first 0.5 sec, and the shortening amount decreased slightly from 0.025 to 0.02 mm over a period of 4.5 sec, which corresponds to the onset of extrusion of molten run materials from the slip zone (Fig. 10b). Significant shortening occurred continuously from 5 sec after the onset of

high-velocity slip (Fig. 10a), which corresponds to the period of extrusion of large amount of molten run products from the slip zone (Fig. 4). In contrast, during the LVF experiments, shortening occurred after 61 sec had passed since the initiation of slip (Fig. 10c–d).

3.5. FTIR and TG–DSC Analyses

To quantitatively analyze the water contents of the starting rock sample and the run products, we performed Fourier transform infrared (FTIR) and thermogravimetry (TG)–differential scanning calorimetry (DSC) spectroscopic analyses using Jasco FT/IR-300E and Netzsch STA 449C Jupiter apparatuses, respectively.

The run products of three HVF experiments (slip rate of 1.10 m/s and normal stresses of 2, 5, and 9 MPa, respectively) were selected for FTIR analysis, which was performed along profiles across the slip zone in thin sections of ~0.25 mm thick. In the FTIR spectra, absorption bands are observed at the OH stretching region (3500–3700 cm⁻¹) in the host serpentine sample but not in the slip zone (‘run product’ in Fig. 11). This result indicates that little water (OH) remained in the slip zone. The water contents were calculated using OH peaks at 3500–3700 cm⁻¹ based on the method proposed by Beran and Libowitzky [2006]. The calculated water contents are shown in Fig. 12.

The FTIR analyses show that no water occurs in the slip zone or in the adjacent dehydration zones, and the water content ranges from 0.5 to 15 wt.% in the host rocks (Fig. 12). These results indicate that all of the water in the slip zone and adjacent dehydration zones (see Fig. 11) was released during the HVF experiments.

278 Four run products, of two HVF (1.10 m/s, No.1 and No.3) and two LVF (0.35 m/s,
279 MH001) experiments performed under normal stresses of 2.4–2.5 MPa, were powdered
280 for TG–DSC analysis. For comparison, one sample of host rock was also analyzed.
281 Approximately 30 mg of sample was placed in a covered Pt₉₀Rh₁₀ crucible and heated
282 from 25°C to 1050°C at a rate of 10°C min⁻¹ under a flow of nitrogen gas (50 ml min⁻¹).

283 The DSC–TG spectra of the host rock sample reveal (i) a weight loss after heating of
284 6.9 wt.%, as determined by TG (Fig. 13a); and (ii) the heat flux during heating showed
285 an endothermic peak at 550–700°C and an exothermic peak at 800–850°C (Fig. 13b).
286 The first peak was accompanied by a weight loss, indicating that it corresponds to a
287 dehydroxylation reaction. The second peak was not accompanied by a weight loss, and
288 its relation to the reactions is currently unknown. The host sample was heated a second
289 time, from 25°C to 1050°C, showing no weight loss and no endothermic peak,
290 confirming that the dehydroxylation was completed during the first heating. The EPMA
291 analyses indicate marked differences in the analytical totals among the serpentine
292 minerals of the host rock (Table 2), the molten material remaining in the slip zone
293 (Table 3), and the dehydration zones (Table 4). The totals for the molten materials and
294 dehydration zones that abut the slip zone are up to 97–99.5 wt.%, about 9–11 wt.%
295 higher than that of the serpentine minerals of the host rock. Considering the presence of
296 ~35% olivine in the host rock sample, the differential amounts of 9–11 wt.% are
297 comparable to the weight loss of 6.9 wt.% in the run products, including molten
298 materials and fine-grained clasts of the host rock, as determined by TG (Fig. 13a).
299 Accordingly, the difference in analytical total reflects the release of water from the
300 molten materials and from serpentine minerals that abut the slip zone, due to the

dehydration of serpentine minerals.

Figure 13 shows the TG –DSC spectra of four run products, revealing that the total weight losses of water are 0.0 wt.% for samples No. 1 and No. 3, 7.3 wt.% for MH001, and 4.3 wt.% for MH002 (Fig. 14a). Samples No. 1 and No. 3 show no significant endothermic peak at around 550–700°C (Fig. 14b), indicating that the dehydroxylation reaction was completed in the HVF samples. In contrast, MH001 and MH002 show a weight loss and an endothermic peak at around 550–700°C (Fig. 14b), similarly to the host rock (Fig. 13b). This result suggests that the dehydroxylation reaction was not completed during the LVF friction experiment at a low velocity of 0.35 m/s.

4. Discussion and Conclusions

4.1. Coseismic Melting and Dehydration Reaction

The ejection of water vapor and the generation of vesicles within run products during the HVF experiments provide conclusive proof that water was released from the slip zone. The flow structures, the vitreous luster of run products (similarly to that of fresh lava), and X-ray diffraction patterns indicate the coexistence of abundant fluid derived from serpentine and frictional melt resulting from frictional heating in the slip zone. The absence of antigorite crystals in the X-ray diffraction spectra of run products can be interpreted as a result that antigorite crystals were completely broken down and/or melted by friction heating.

As shown in Fig. 9, the molten materials have lower SiO₂ and Al₂O₃ components and

324 higher FeO component similar to that of olivine mineral contained in the host rock than
325 that of serpentine mineral in the dehydration zone, and higher CaO and FeO components
326 than that of serpentine contained in the host serpentinite. The FeO component of the
327 molten materials is about 5~7 wt% higher than that of serpentine minerals contained in
328 the host serpentinite and the dehydration zones, which is similar to that of olivine
329 mineral contained in the host rock (Fig. 9). Furthermore, MgO and Al₂O₃ components of
330 the molten materials are locally similar to that of olivine mineral, higher than that of the host
331 serpentine minerals contained in the host serpentinite and the dehydration zones (Fig. 9).
332 These textural and petrologic features including the chemical compositions of molten
333 materials reveal that not only the serpentine but also the olivine minerals were
334 melted by friction heating, and that a high temperature reached the bounding zones of
335 the slip zone at which serpentine was involved in a dehydration reaction. Previous
336 studies show that the high-velocity friction melting occurs in a chemically
337 non-equilibrium process (Lin and Shimamoto, 1998; Lin, 2008) and that the melting or
338 breakdown of serpentine mineral occurs at temperature of 400~600 °C [Spray, 1992].
339 Thus, the serpentine and olivine minerals are considered to be melted at temperature of
340 \geq melting points of these minerals. Based on the melting point (~1890 °C) of olivine
341 [Spray, 1992], it is estimated that the temperature in the slip zone reached at least up to
342 1890 °C during the HVF experiments.

343 Previous studies indicate that the dehydration of serpentine occurs at temperature of
344 $\geq 500^{\circ}\text{C}$ [Ulmer and Trommsdorff, 1995; Viti, 2010] to 856°C [Lange and Ahrens,
345 1982]. In previous HVF experiments using marble samples, direct thermometer-based
346 measurements of temperature indicate that the slip-zone temperature can reach

600–900°C during the first 0.1 sec of the experiment at a slip rate of 1.17 m/s [Han et al., 2007]. TG, DTG and differential thermal analyzer (DTA) analyses show that serpentine dehydration takes place between 550 and 800°C [Viti, 2010]. Calculation results for an HVF experiment on serpentinite at a slip rate of 1.1 m/s reveal that the flash temperature can rise by ~1250°C over a period of <0.1 sec [Hirose and Bystricky, 2007] and that the temperature can reach 550–1100°C within a narrow slip zone (<3 mm wide) over a period of 2–3 sec [Fukuchi and Imai, 2001]. Furthermore, seismic data reveal that if the shear energy is constrained within 1 cm of the slip plane during seismic slip, the temperature can readily increase to 100–1000°C during a medium- to large-magnitude earthquake [Kanamori and Brodsky, 2004]. Thus, the temperature up to 1890°C in the present HVF experiments as documented above, within a dehydration zone of <2 mm in width (Fig. 12) and at slip rates of 1.10–1.31 m/s, is sufficient to induce the dehydration of serpentine.

The breakdown of serpentine to olivine + talc + water and the breakdown of talc to enstatite + quartz + water at 1 bar follow a second-order rate law [Bose and Ganguly, 1994]. The breakdown of antigorite in the present experiments is indicated by the consumption of antigorite and increased integrated intensities of olivine in the X-ray diffraction spectra for the run products, although talc could not be identified from the spectra. Direct observations made during the present experiments also demonstrate that serpentine dehydration reaction occurred within 0.2 sec of the onset of slip. Furthermore, the antigorite crystals observed in the dehydration zones that abut the slip zone show textural features indicative of decomposition, without the distinct cracks and fibrous texture observed in the starting sample of host rock (Figs 5 and 6a–c). This finding

demonstrates that the serpentine dehydration reaction that occurred in a wide zone involving the slip zone was caused mainly by the diffusion of frictional heat from the slip zone. Our results support the idea that the diffusion of frictional heat is a major coupling factor of thermal pressurization caused by the dehydration of serpentine within seismogenic fault zone [Brantut et al., 2010].

4.2. Mechanism of Dynamic Weakening

The dynamic weakening of faults during earthquakes is considered to result from several mechanisms, including friction melting [Spray, 1993; Hirose and Shimamoto, 2005; Di Toro et al., 2006, 2011], the presence of S–C cataclasites [Lin, 1999, 2008], the formation of fault gouge [e.g., Rice, 2006; Mizoguchi et al., 2007], thermal pressurization [e.g., Lachenbruch and Sass, 1980; Wibberley and Shimamoto, 2005], dehydration reactions involving water-rich minerals such as serpentine and kaolinite [e.g., Hirose and Bystricky, 2007; Brantut et al., 2008; Viti and Hirose, 2010; Kohli et al., 2011] as well as fault gouges present in seismogenic fault zones [e.g., Famin et al., 2008; Hamada et al., 2009, 2011], thermal decomposition [Han et al., 2007], and the dehydroxylation of mica minerals [Hirono and Tanikawa, 2011]. Each of these mechanisms is closely linked to frictional heating generated upon the fault plane during seismic slip. The frictional melt generated in the slip zone also plays an important role in lubricating the slip plane, thereby facilitating further slip along seismogenic faults during large earthquakes [e.g., Di Toro et al., 2006, 2011]. Although dynamic fault strength during frictional melting is commonly considered to be controlled by the

development of a molten layer and by the viscosity of this layer [Spray, 1993; Hirose and Shimamoto, 2005], the serpentinite dehydration may also be a main mechanism resulting in a dramatic increase in pore pressure that may lead to the dynamic weakening of seismogenic faults. The outspray of water vapor within the first 0.2 sec of each run as observed during the HVF experiments indicates that the dehydration of serpentine occurred in the first 0.2 sec. The fact that molten materials were squeezed out after 0.5 sec after the onset of the slip from the slip zone as observed during the experiments reveals that a molten layer was produced in the same time as the water vapor within the slip zone during the first 0.5 sec. These findings indicate that both water vapor and molten material were co-existed in the slip zone in the first 0.5 sec of the run. It may be difficult to produce sufficient fault gouge material to affect the fault strength over a short duration of 0.2 sec because of small amount of generated slip. Accordingly, the dramatic decrease in the frictional coefficient from 0.56 to 0.25 in the first 0.2 sec after the onset of slip (Fig. 10a–b) can be interpreted as a result of thermal pressurization caused by the dehydration of serpentine that accompanies frictional melting during the HVF experiments. In contrast, similar weakening was observed in the LVF experiments, in which the frictional coefficient decreased from 0.55 to 0.25, but this took ~10 sec after the onset of run (Fig. 10c–d). This may be interpreted as that enough amount of fine-grained material (fault gouge) formed in the slip zone which weakened the fault strength as proposed by previous studies [Rice, 2006; Mizoguchi et al., 2007].

In the present study, it is considered that the main weakening mechanism of the simulated faults under conditions equivalent to large earthquake slip was the thermal pressurization caused by serpentine dehydration associated with frictional heating and

416 melting, resulting in the excess saturation pressure of water vapor and fluid including
417 melt in the slip zone. Previous studies show that frictional heating would cause the fluid
418 in the fault gouge to expand in volume much more than would be the solid cage [e.g.,
419 Sibson, 1973, Lachenbruch and Sass, 1980]. Thus, shear-induced dilatancy of the gouge
420 overwhelms the thermal expansion effect and a pressure increased must be induced in
421 the pore pressure [Rice and Cocco, 2007]. The saturation pressure of water vapor is
422 known to vary with the temperature of the vapor [Kaye and Laby, 1986]: when confined
423 water is heated, the saturation pressure of water vapor (P) shows a rapid increase.
424 Assuming ideal gas (water vapor) behavior, $P = nRT/V$, where n is the number of moles,
425 R is the water vapor constant, T is absolute temperature (K), and V is the volume of
426 water vapor. From room temperature (20°C) to 360°C, P increases by three orders of
427 magnitude, from $\sim 10^{-2}$ to ~ 20 MPa [Kaye and Laby, 1986]. Theoretically, under the
428 pressure–temperature conditions of seismogenic depths, the thermal expansivity of
429 water is on the order of $10^{-3}/^{\circ}\text{C}$, meaning that the increase in fluid pressure at the depths
430 of hypocenters would be on the order of 1 MPa/ $^{\circ}\text{C}$ [Lachenbruch and Sass, 1980;
431 Kanamori and Brodsky, 2004]. Thus, the corresponding pressure of water vapor and
432 fluid would be >20 MPa in a simulated fault zone under conditions equivalent to large
433 earthquake slip, comparable to the stress drop that occurs during large earthquakes
434 [O’Hara et al., 2006]. In the present HVF experiments, the decrease in the average
435 shortening amount in the first 5 sec after the onset of slip can be interpreted as a result
436 of pore pressure increase due to the thermal pressurization that accompanies frictional
437 melting within the slip zone. The subsequent increase in the average shortening amount
438 could reflect the extrusion of run products (including vapor, melt, and fine-grained

materials) from the slip zone, as shown in Fig. 4.

Our results support the idea that the thermal pressurization of fluid released by the dehydration of serpentine and friction melt plays an important role in the dynamic weakening of faults during seismic slip within seismogenic fault zones.

4.3. Tectonic Implications

The present experimental results have significant implications for subduction zone earthquakes and transform faults such as the San Andreas Fault. Seismogenic faults within a subduction zone are generally associated with a mantle wedge composed of serpentinites that contain ~13 wt.% H₂O [Ulmer and Trommsdorff, 1995]. Hydrous antigorite may play an important role in the transport of water into the earth when it enters a subduction zone. The present results show that the frictional heat generated upon a seismogenic slip plane under conditions of large earthquake slip can lead to melting in a seismic slip zone and the dehydration of serpentinite within a wide dehydration reaction zone (comprising the main slip zone and bounding zones) due to the diffusion of frictional heat from the slip plane, thereby increasing the pore pressure due to the thermal expansivity of fluid and water vapor. Furthermore, the frictional melt generated in the slip zone also plays an important role in lubricating the slip plane, facilitating further slip along seismogenic faults during large earthquakes (Di Toro et al., 2011). Accordingly, our experiment results confirm that the serpentine dehydration that accompanies frictional melting results in a sudden increase in pore pressure, which in turn may lead to a reduction in the effective normal stress across the fault, in turn

resulting in a marked reduction in the dynamic fault strength, thereby enabling further slip along the fault in a subduction zone.

This sequence of events might also influence the physics of earthquake processes within subduction zones, explaining the large seismic slip of >15–50 m generated by huge subduction-zone earthquakes such as the 2004 M_w 9.4 Sumatra–Andaman earthquake [Subarya et al., 2006] and the 2011 M_w 9.1 Tohoku Japan earthquake [Yagi and Fukahata, 2011]. In addition, if the large volumes of fluid released by the dehydration of serpentinites during huge earthquakes are concentrated upon and migrate along pre-existing fault zones to neighboring fault segments, such a sudden increase in fluid pressure would promote further brittle failure in barrier areas of seismogenic faults and in neighboring fault segments. The present results indicate that the dynamic parameters of seismogenic faults are strongly affected by serpentine dehydration during large earthquakes within subduction zones and along intracontinental faults that contain abundant hydrous minerals.

Acknowledgments

We are grateful to Prof. T. Shimamoto for designing the rotary-shear low- to high-velocity test machine. This work was supported by a Grand-in-Aid for Scientific Research (Science Project No. 23253002 for A. Lin) from the Ministry of Education, Culture, Sports, Science and Technology of Japan.

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611 Captions to figures

612 Figure 1. Diagram showing the rotary-shear low- to high-velocity frictional testing
613 machine equipped with a hydraulic pressure vessel. 1: specimen holder (one-axis
614 rotary-shear specimen box); 2: pressure vessel for elevated pore-water pressure;
615 3–4: torque and axial force gauges; 4: electromagnetic clutch; 5: force gouge; 6:
616 oil pressure cylinder; 7: servo-motor and gear system; 8: motor controller; 9:
617 manual pump; 10: oil pressure pump; 11–12: manual oil pressure pump; 13:
618 pressure generator system; 14–15: data recording system. This machine was
619 installed at Shizuoka University, Japan, in early 2007.

620 Figure 2. Close-up view of the specimen holder (a), configuration of the specimen set (b),
621 and orientation of the thin section cut from the run products (c). The thin section is
622 oriented perpendicular to the slip zone.

623 Figure 3. Petrographic features of serpentinite used for the frictional experiments in this
624 study. (a) Photomicrograph (cross polarized light); (b) scanning electron image.
625 The serpentinite is composed of 65% serpentine (antigorite) and 35% olivine.

626 Figure 4. Photographs taken through the transparent window cover of the experimental
627 apparatus showing dust and water vapor (a, c, d), and sparks of yellow–red melt (b)
628 ejected from the interface (slip zone) of two cylindrical serpentinite specimens at a
629 slip rate of 1.10 m/s and normal stress of 6.5 MPa during a high-velocity friction
630 (HVF) experiment (HVF R001). The water vapor cooled to form water droplets
631 (Wd) upon the window cover along a zone of 2–7 mm in width during the 20 sec
632 duration of the experiment.

633 Figure 5. Photomicrograph (a) and back-scattered electron micrograph (b) showing

634 representative microstructures of a simulated fault zone (slip zone) in a
635 high-velocity friction (HVF) experiment (run HVF R001) conducted at a slip rate
636 of 1.10 m/s. Note that dehydrated zones (dz) are developed on both sides of the slip
637 zone.

638 Figure 6. Back-scattered electron micrographs showing representative microstructures
639 of a simulated fault zone (slip zone) in experiment run HVF R001 conducted at a
640 slip rate of 1.10 m/s, showing (a) the slip zone and dehydration zone (dz); (b) a
641 close-up view of the dehydrated zone in serpentine; (c) a close-up view of (b); and
642 (d) representative textures of run products ejected from the slip zone, showing melt
643 textures and flow layers, vesicles (v), and a fine-grained matrix (molten material).

644 Figure 7. Representative textures of run products. (a) Photograph of fine-grained matrix
645 (molten material) ejected from the slip zone, showing flow textures, vesicles, and a
646 vitreous luster similar to that of fresh lava. (b) Back-scattered electron micrograph
647 showing the textures of vesicles and fine-grained clasts (HVF R001 run product).

648 Figure 8. Powder X-ray diffraction spectra derived from (a) the host serpentinite, (b)
649 reference olivine, and (c–f) representative run products remaining in the slip zone,
650 for experiments with a slip rate of 1.31 m/s and normal stress of 5 MPa (run HVF
651 R012) (c), and a slip rate of 1.10 m/s and normal stress of 6.5 MPa (run HVF
652 R001) (d). The sample (HVF 012) was picked from the run product material
653 extruded from the slip zone, and the sample (HVF R001) was separated from the
654 run product remained in the slip zone. (e) Enlargement of (d) between 2 θ values of
655 15° and 35°, and (f) calibrated sample containing 90 wt.% fine-grained clasts of
656 run product and 10 wt.% glass (obsidian). ol: olivine; ant: antigorite. Note that the

run product (HVF R001) showed a distinct, broad range of 2θ values from 20° to 35° , as observed for the calibrated sample containing 10 wt.% glass (f). The vertical axis (diffraction density) is the same for all spectra. CPS: counts per second.

Figure 9. Variation of major oxides in the host serpentinite, slip zone and dehydration zone generated in the HVF experiments. Data are shown in Tables 2-4.

Figure 10. Friction coefficient versus time for experiment run HVF R012 at a slip rate of 1.31 m/s and normal stress of 5 MPa (a–b), and run LVFR110 at a slip rate of 0.35 m/s and normal stress of 8 MPa (c–d). (b) Enlargement of the first 2.4 sec of the HVF experiment shown in (a). (d) Enlargement of the first 20 sec of the LVF experiment shown in (c).

Figure 11. Representative FTIR spectra of the host rock (serpentinite) and run product of a high-velocity friction experiment (MH R0012). Note that the OH peaks are present at $3500\text{--}3700\text{ cm}^{-1}$ in the host serpentine sample but not in the slip zone.

Figure 12. Water contents of the high-velocity friction (HVF) run products calculated from the Fourier transform infrared (FTIR) spectra. Note that the narrow and wide zones reflect a change in slip zone width, as shown in Figs 5 and 6.

Figure 13. Thermogravimetry–differential scanning calorimetry (TG–DSC) spectra of the host rock (serpentinite). (a) TG, (b) DSC. Note that the heat flux during the heating showed an endothermic peak at $550\text{--}700^\circ\text{C}$ and an exothermic peak at $800\text{--}850^\circ\text{C}$.

Figure 14. Thermogravimetry–differential scanning calorimetry (TG–DSC) spectra of the run products of high-velocity friction (HVF) experiments (No. 1 and No. 2),

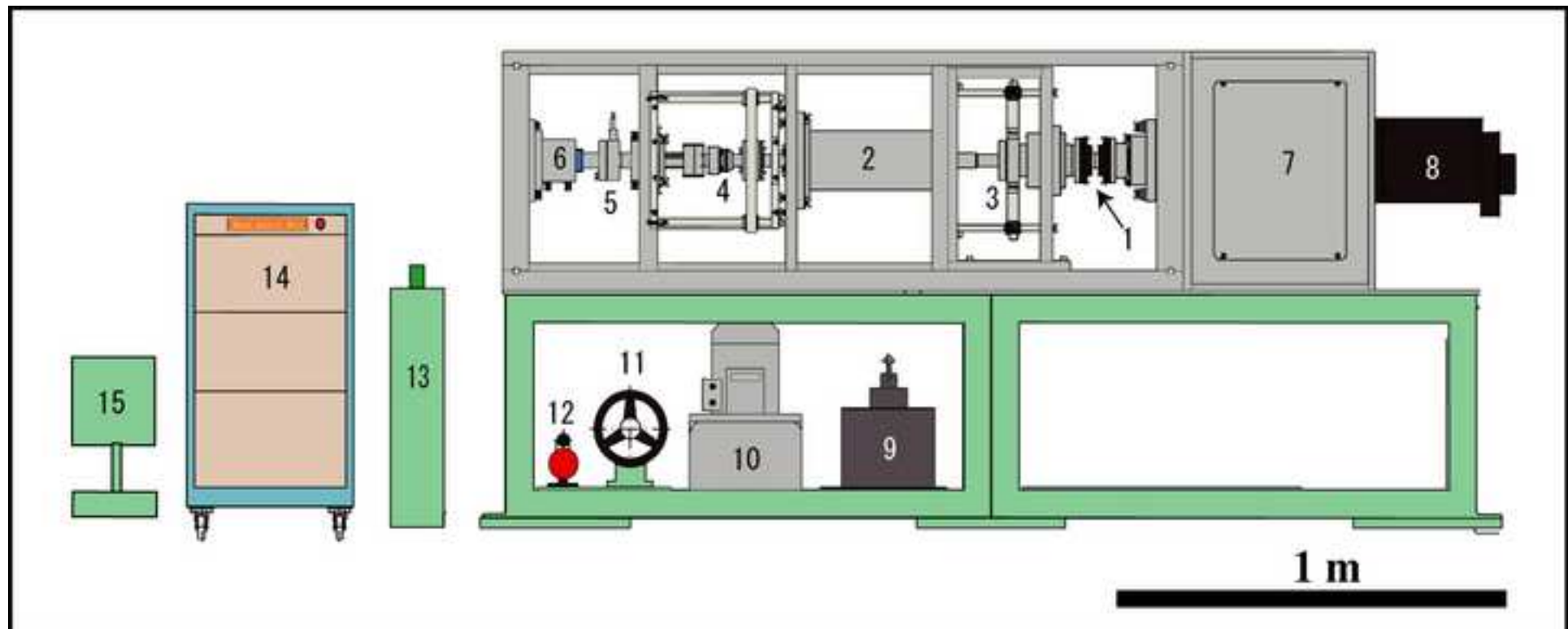
680 and low-velocity friction (LVF) experiments (MH001 and MH002). (a) TG, (b)

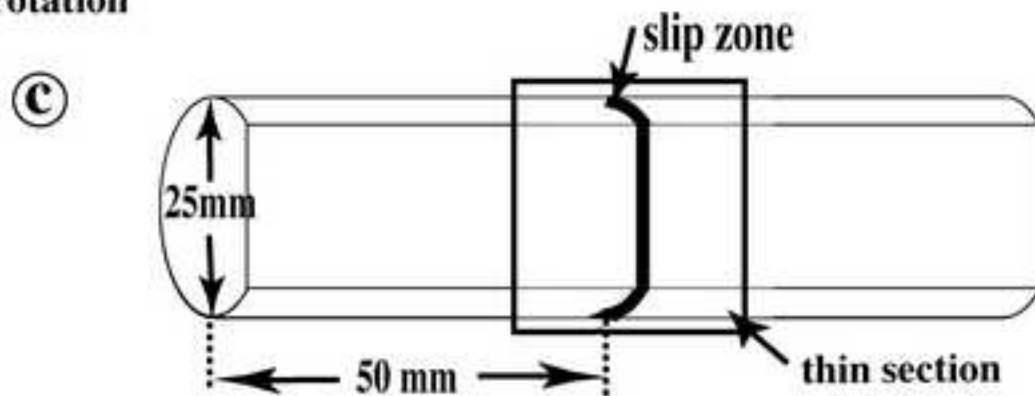
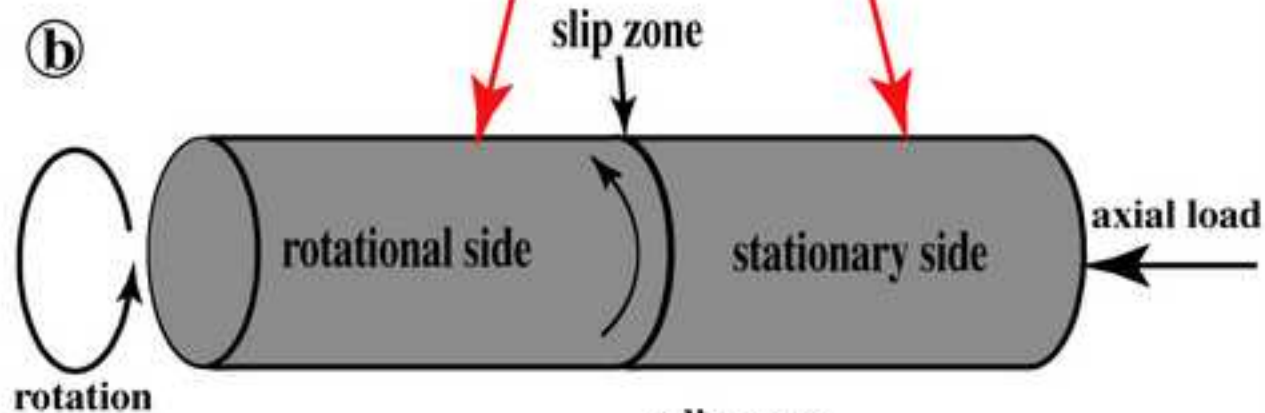
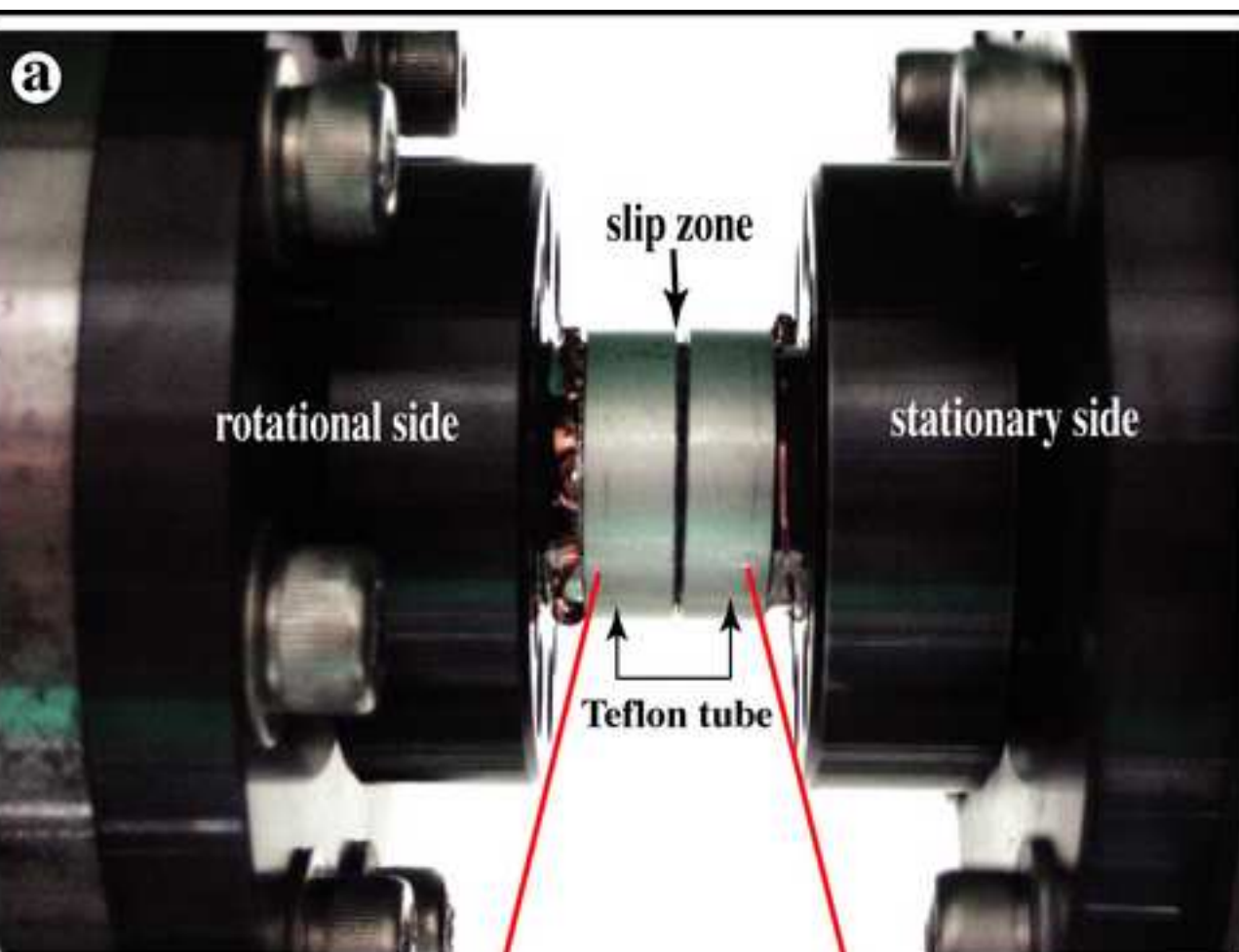
681 DSC. See the text for details.

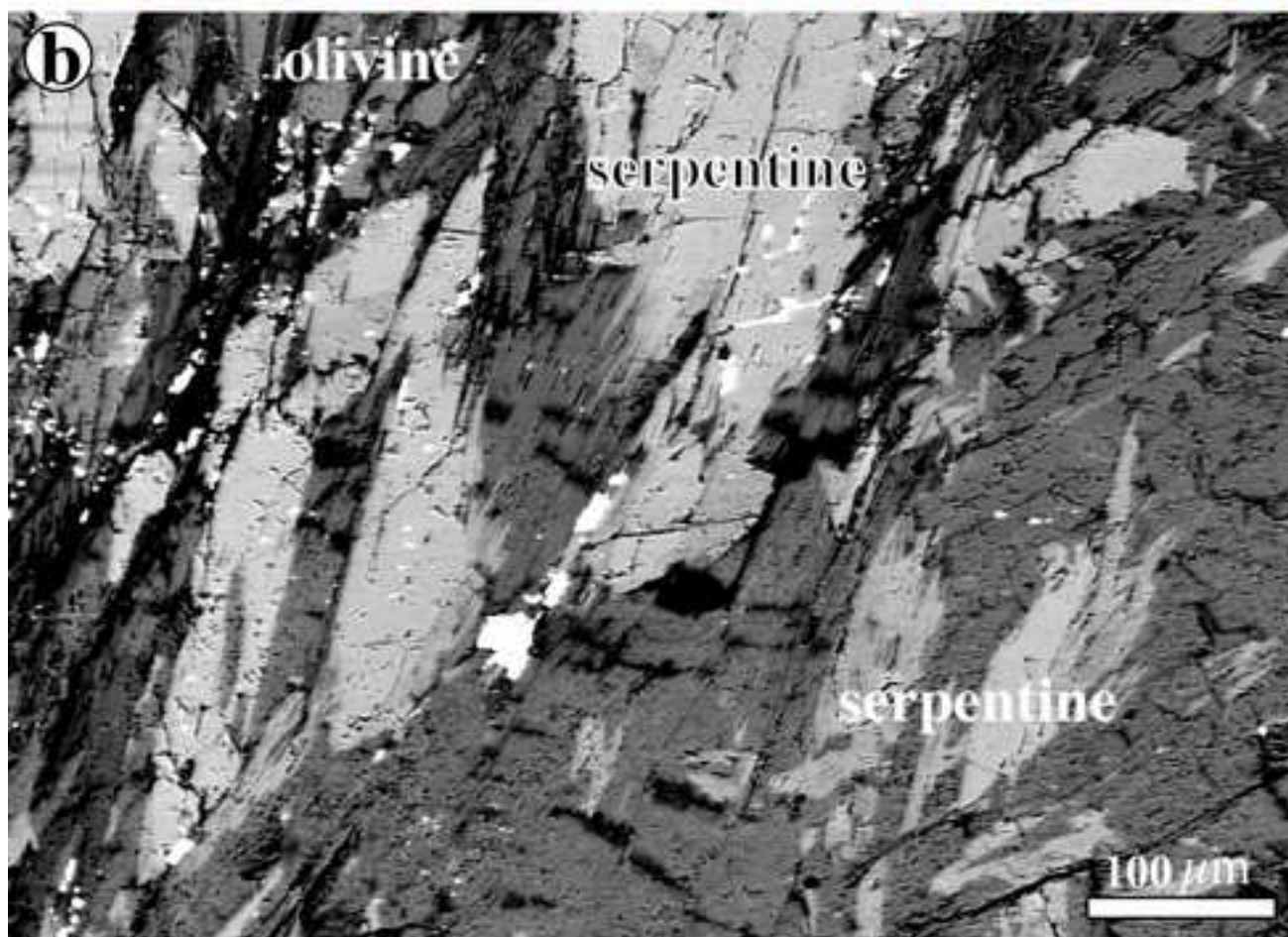
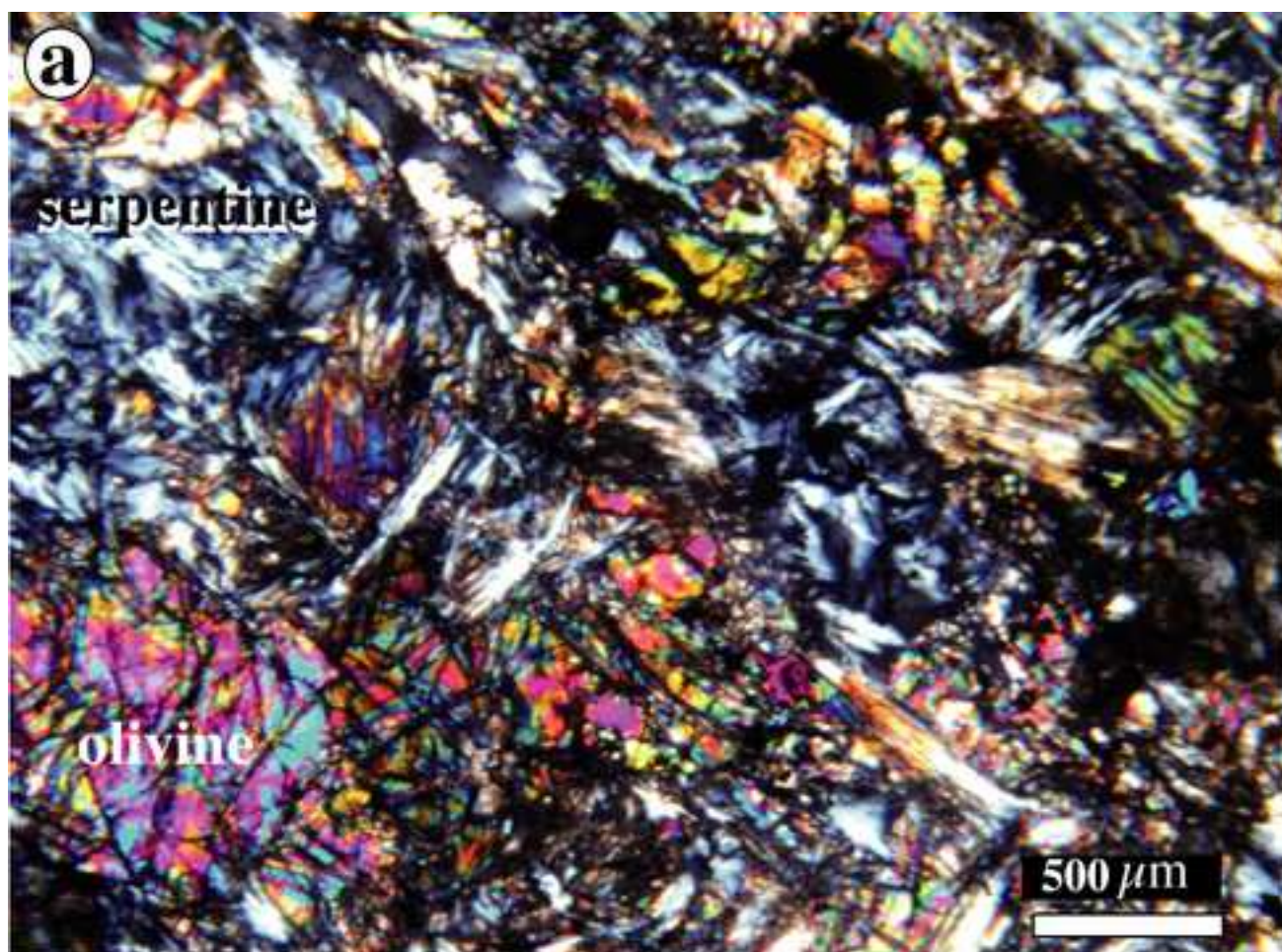
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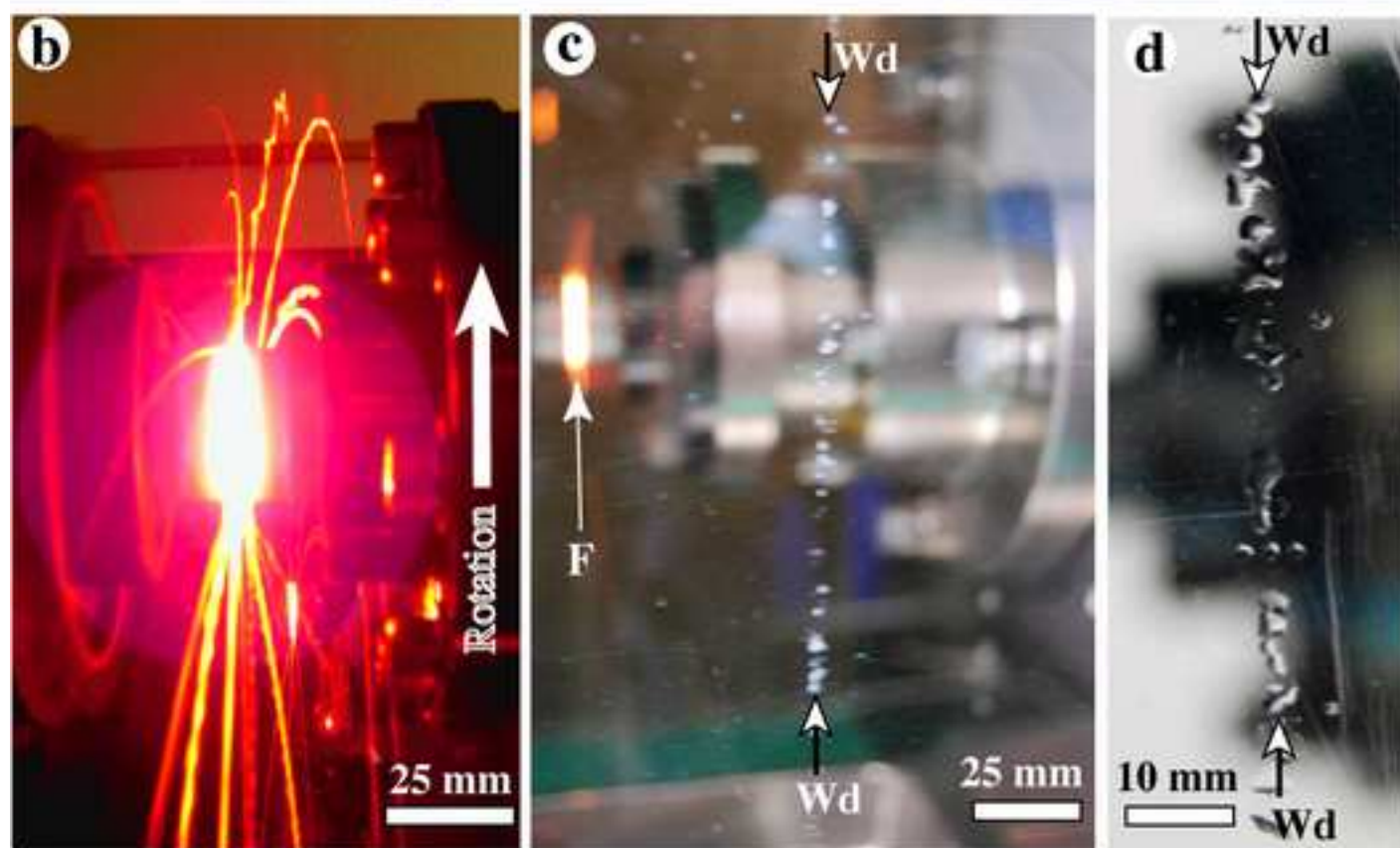
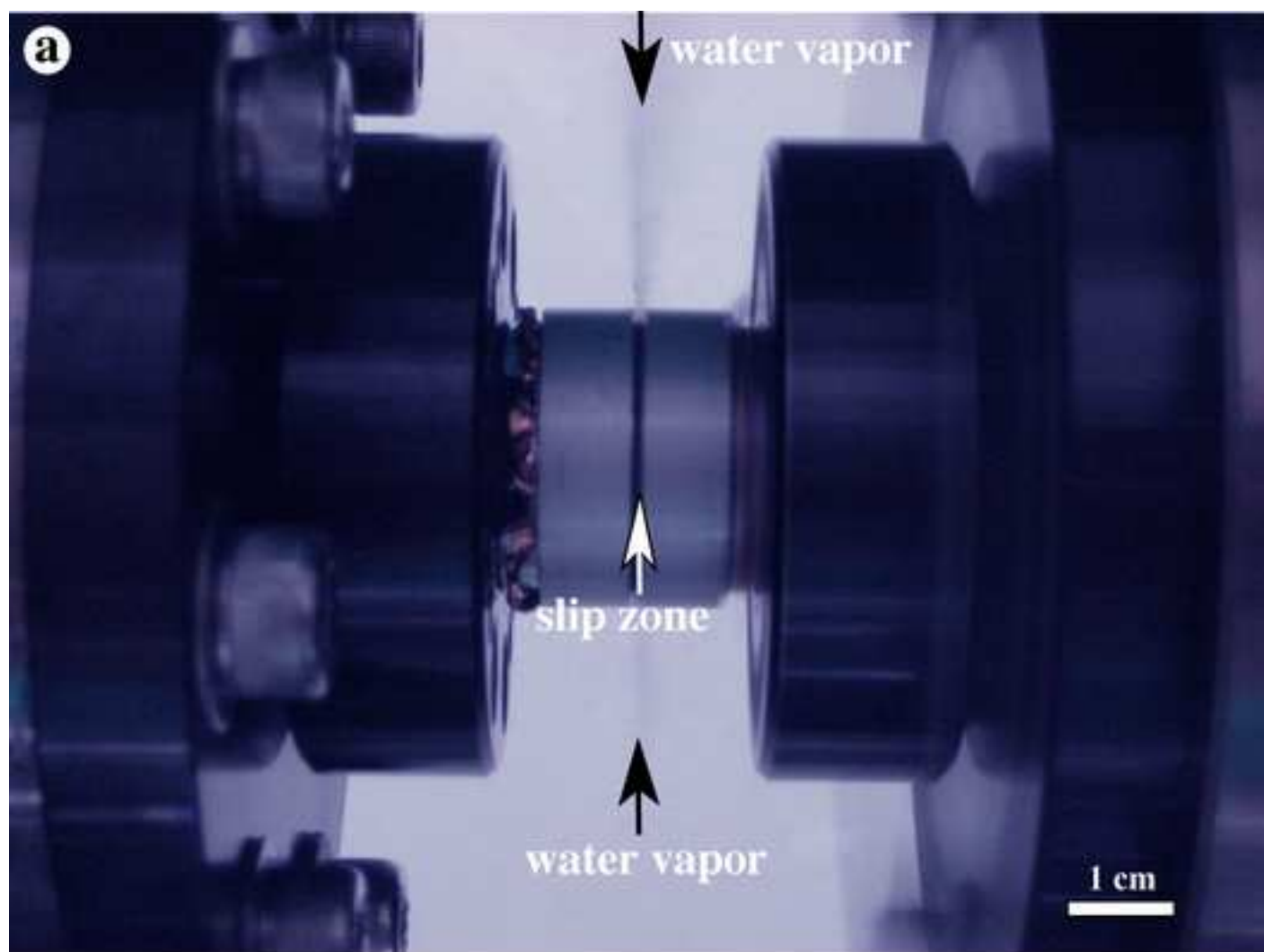
Figure1

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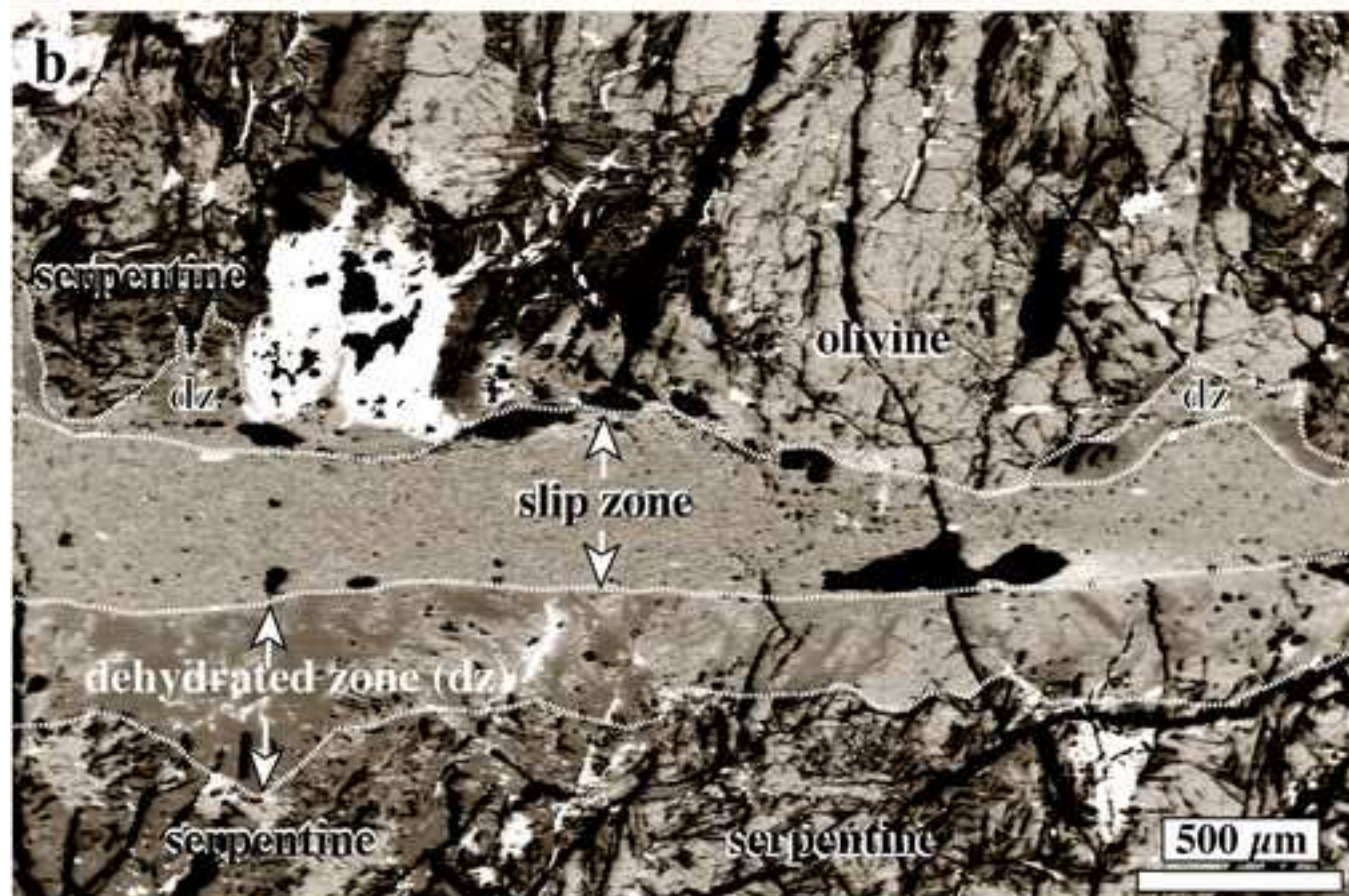
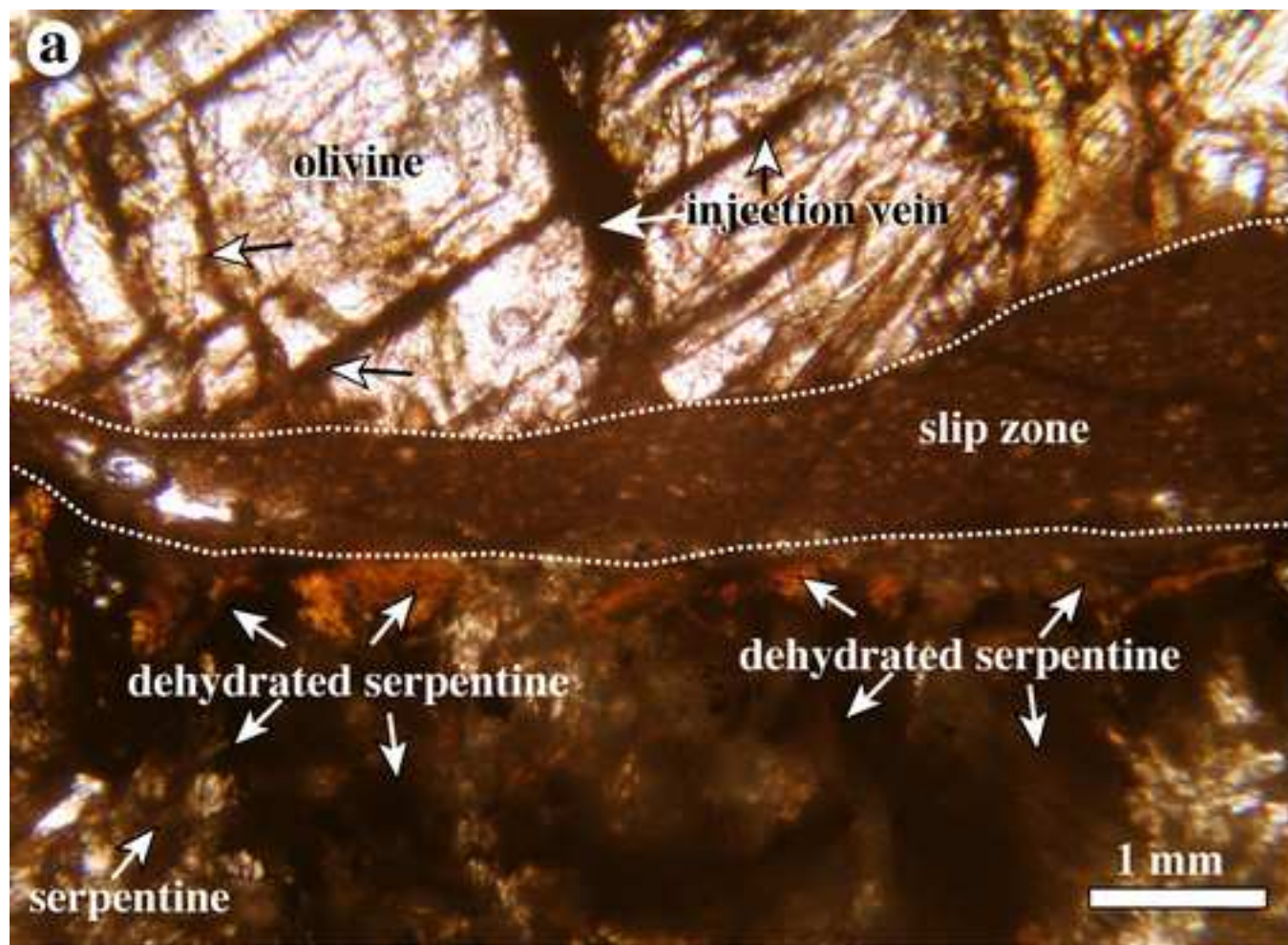
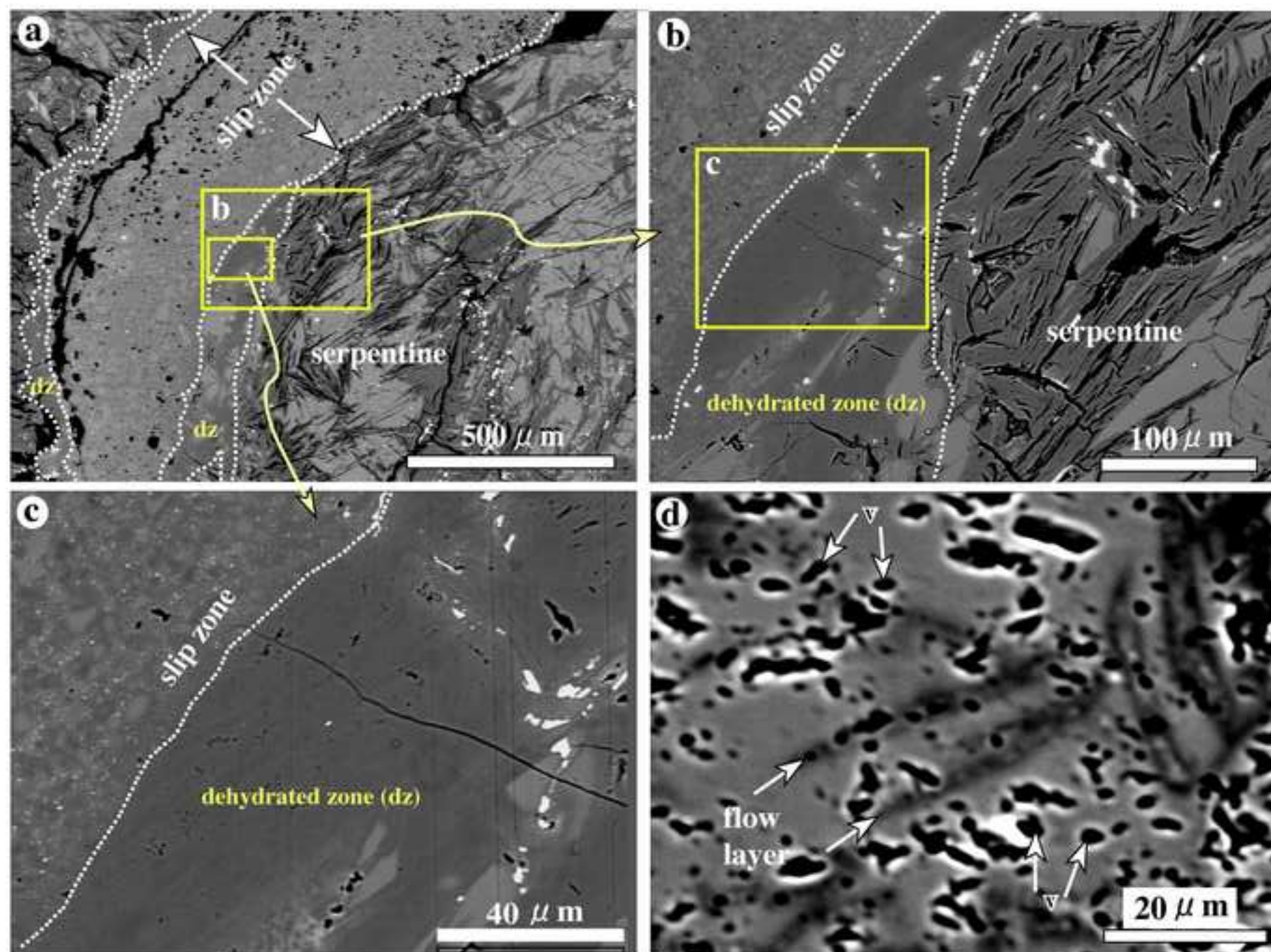
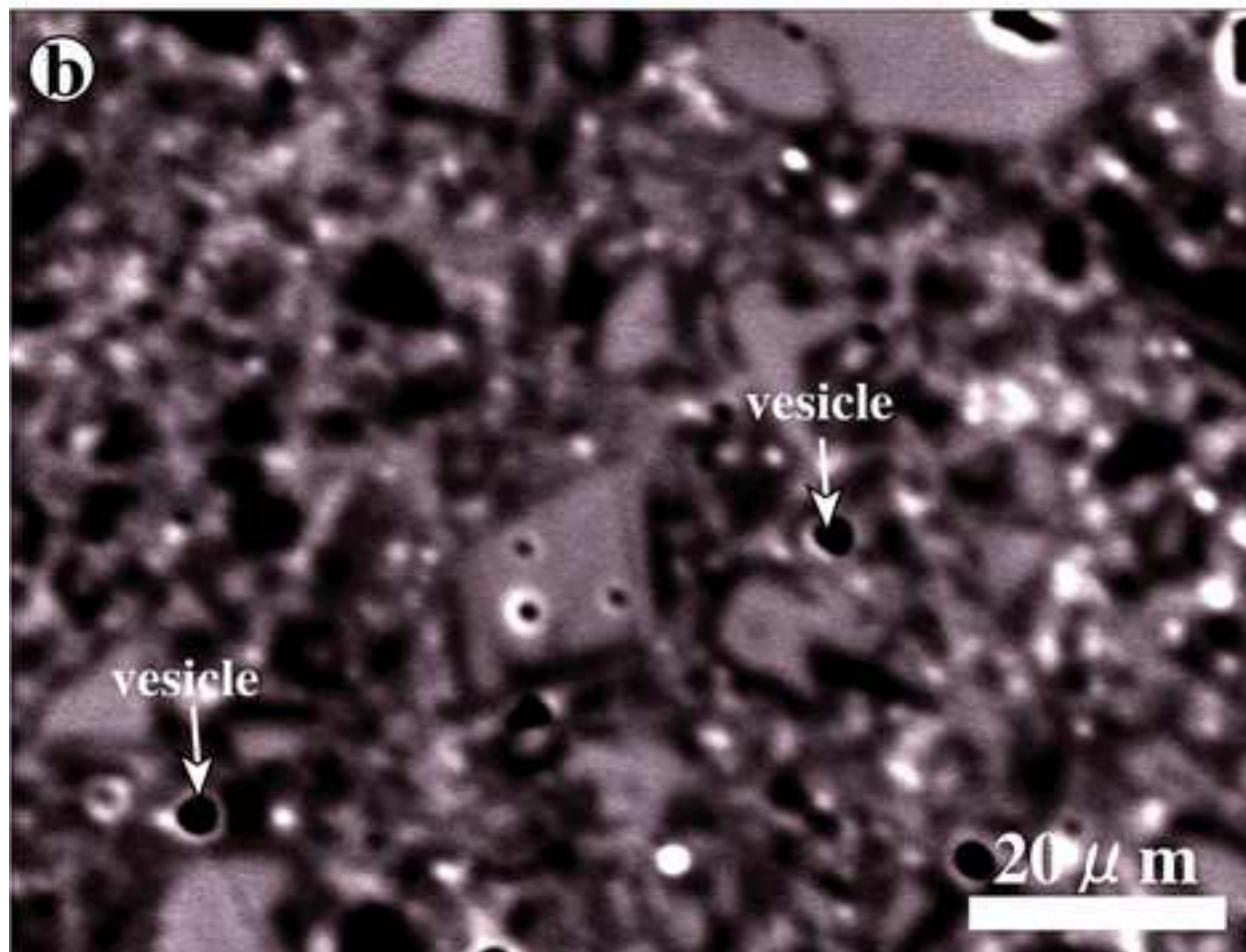
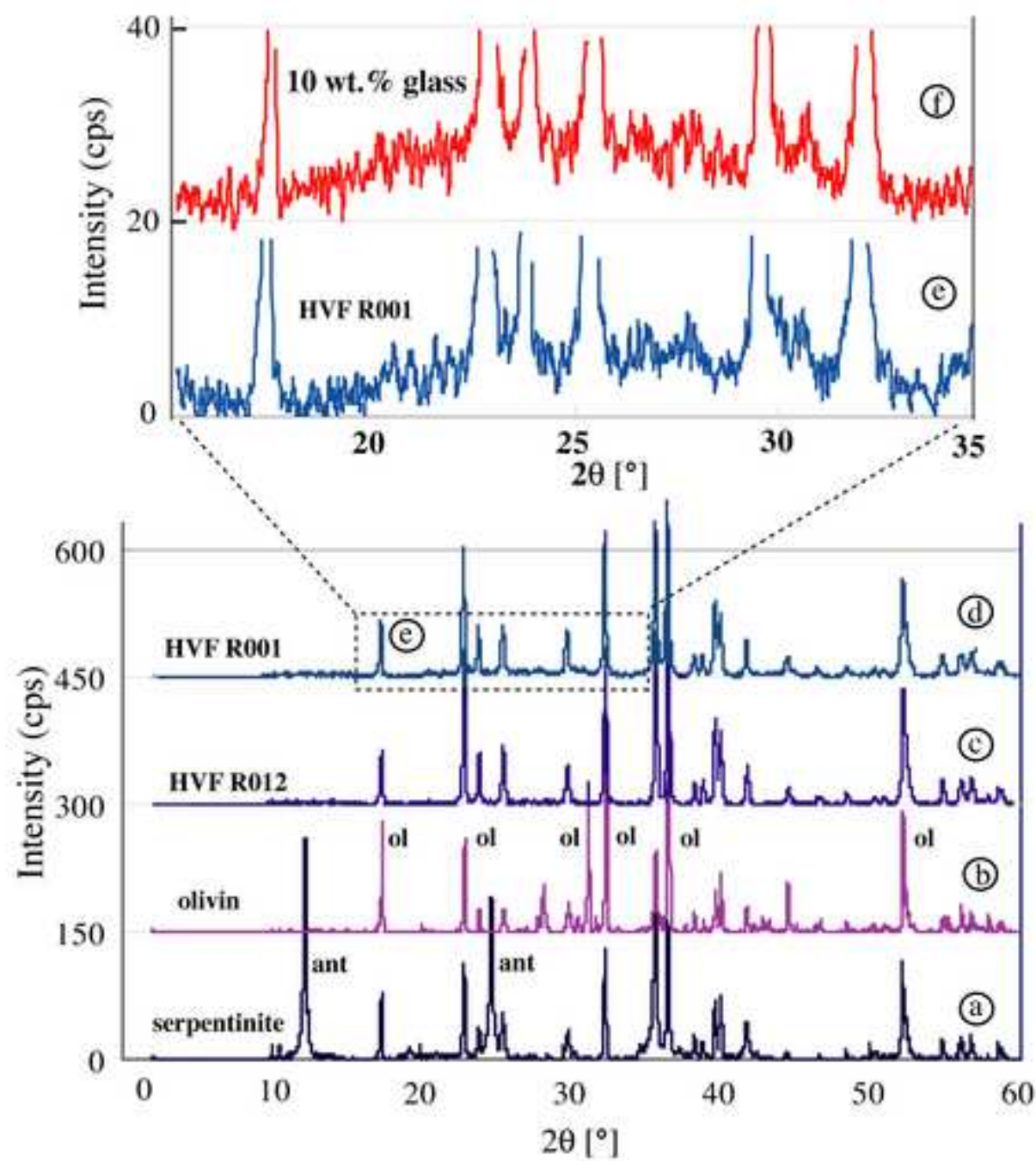


Figure6

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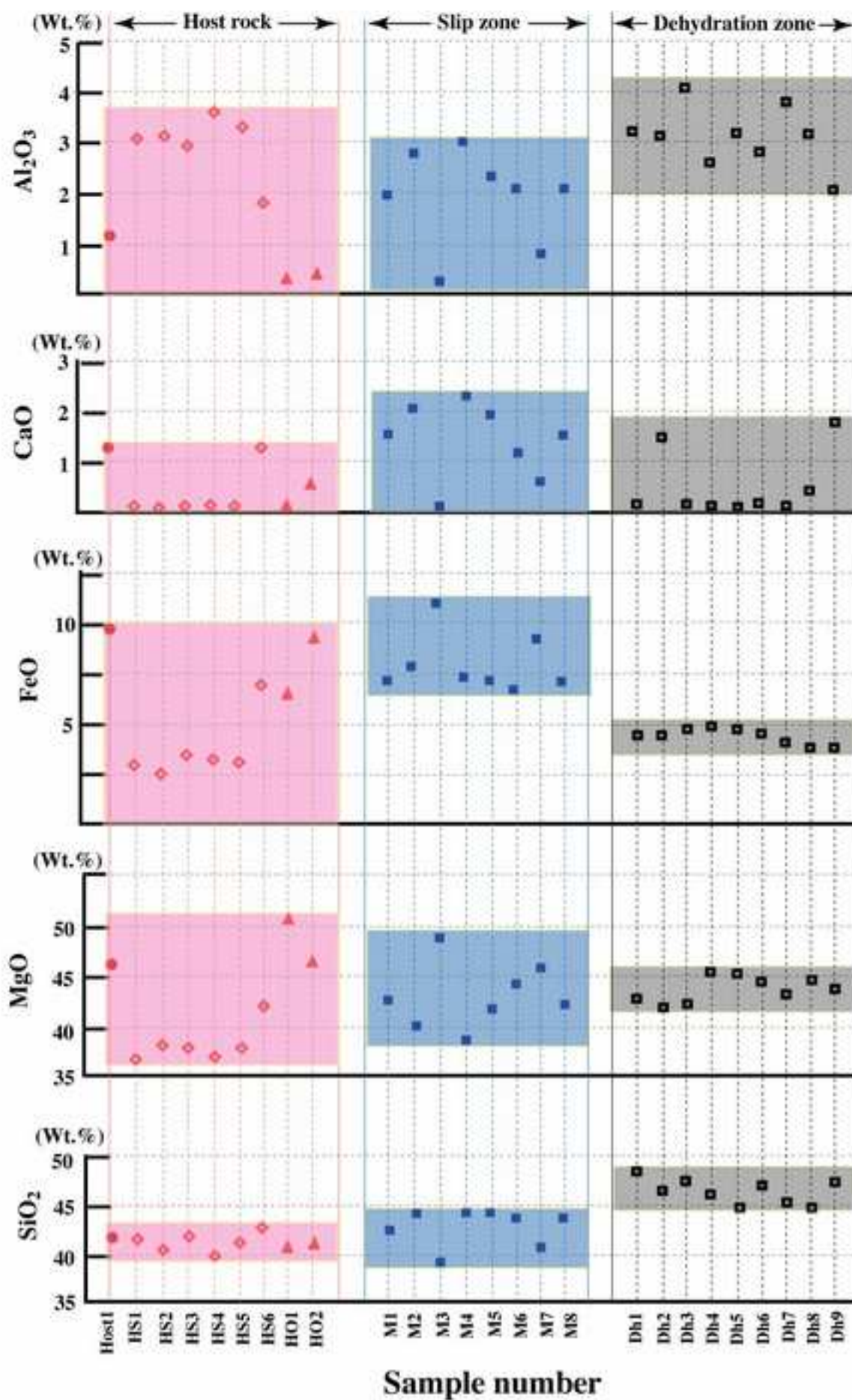


Figure10

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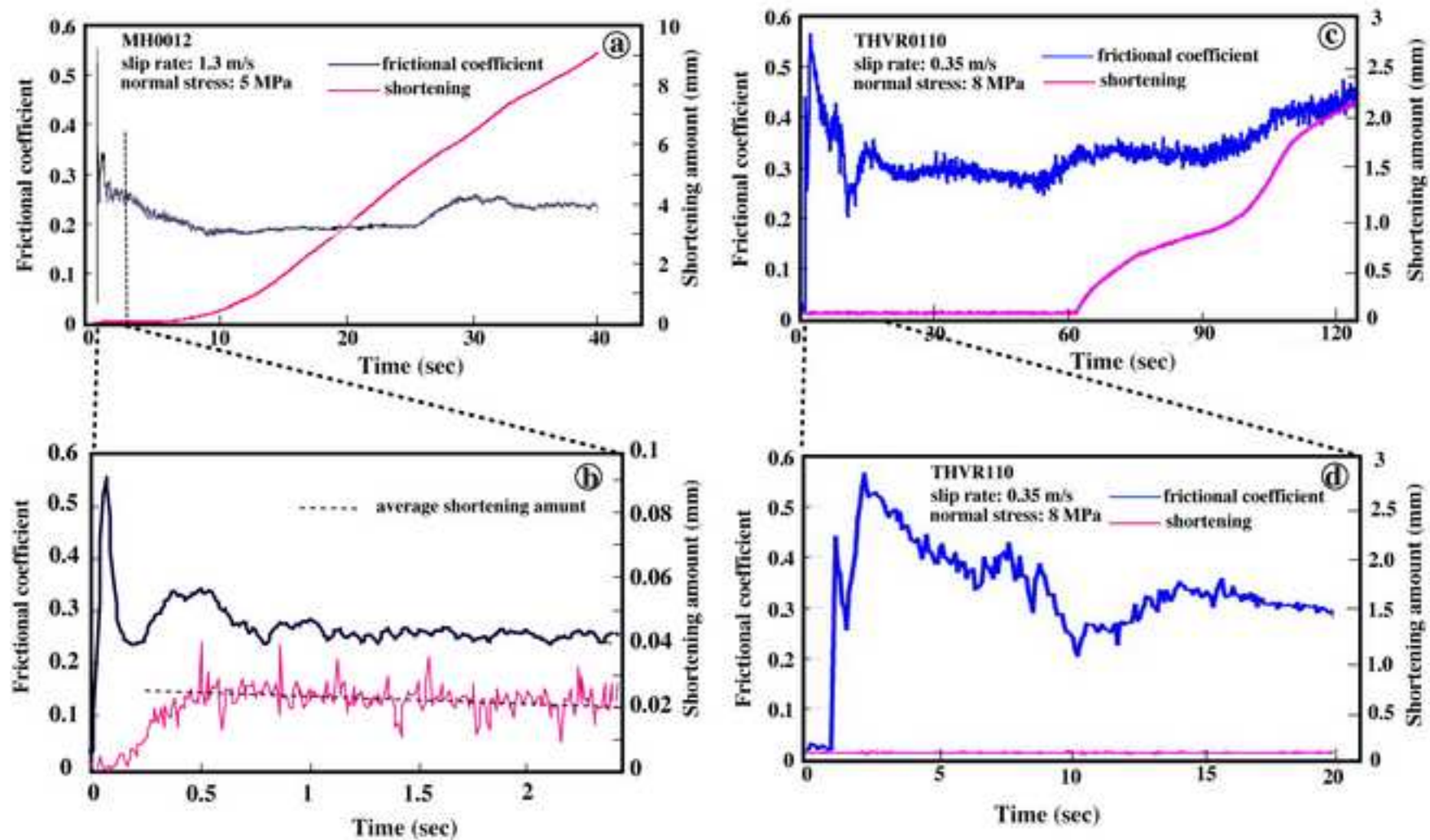


Figure11

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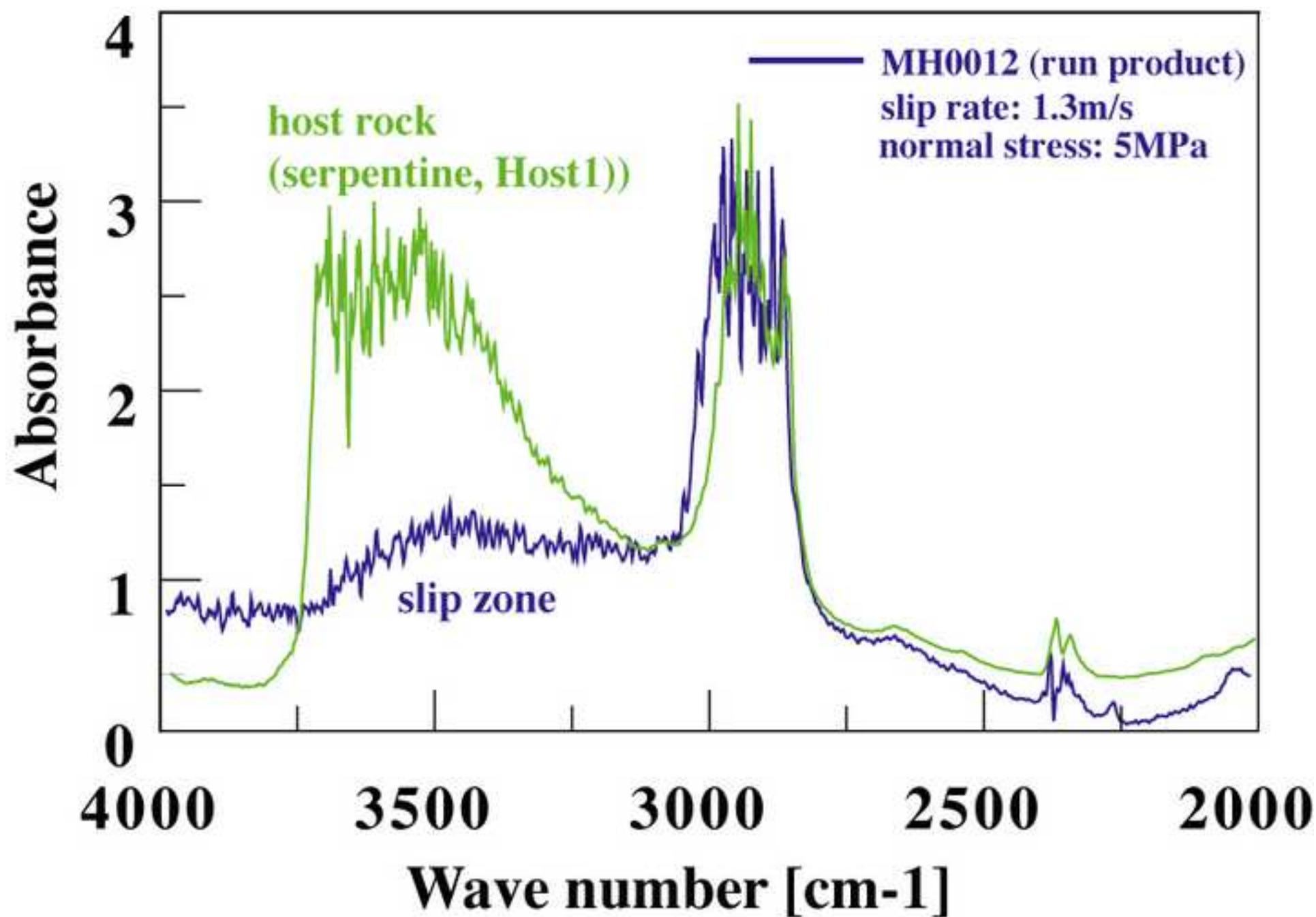
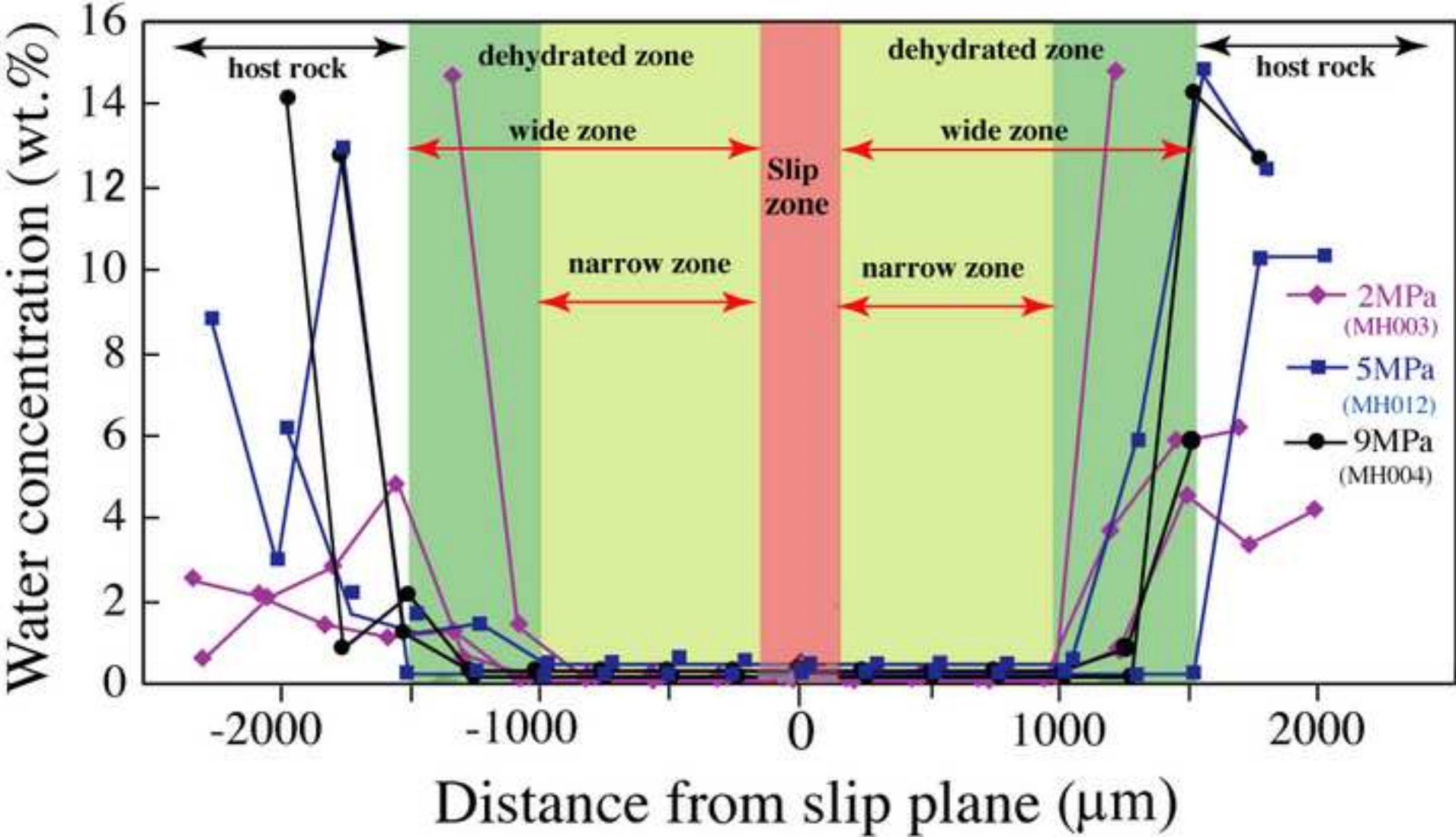
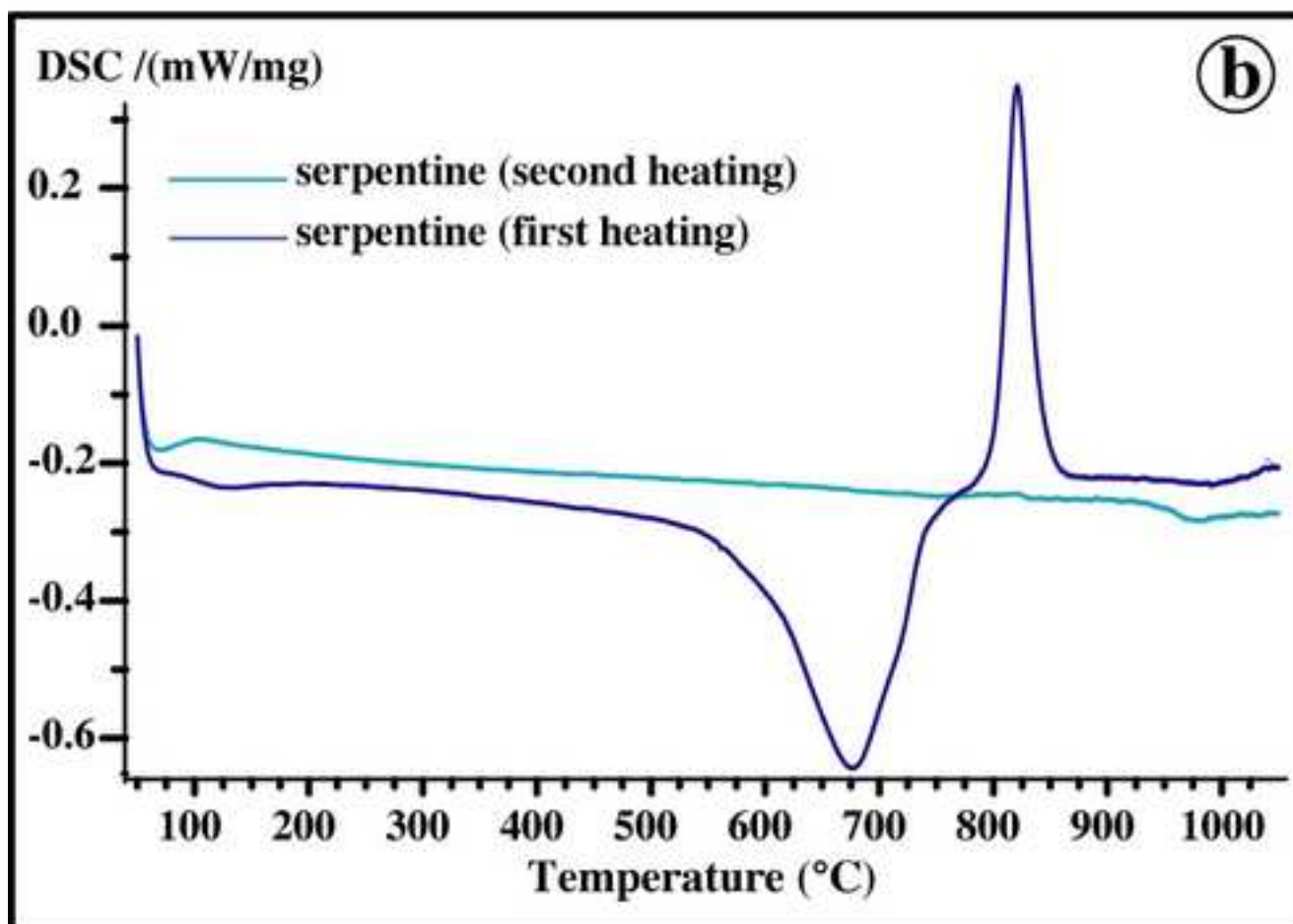
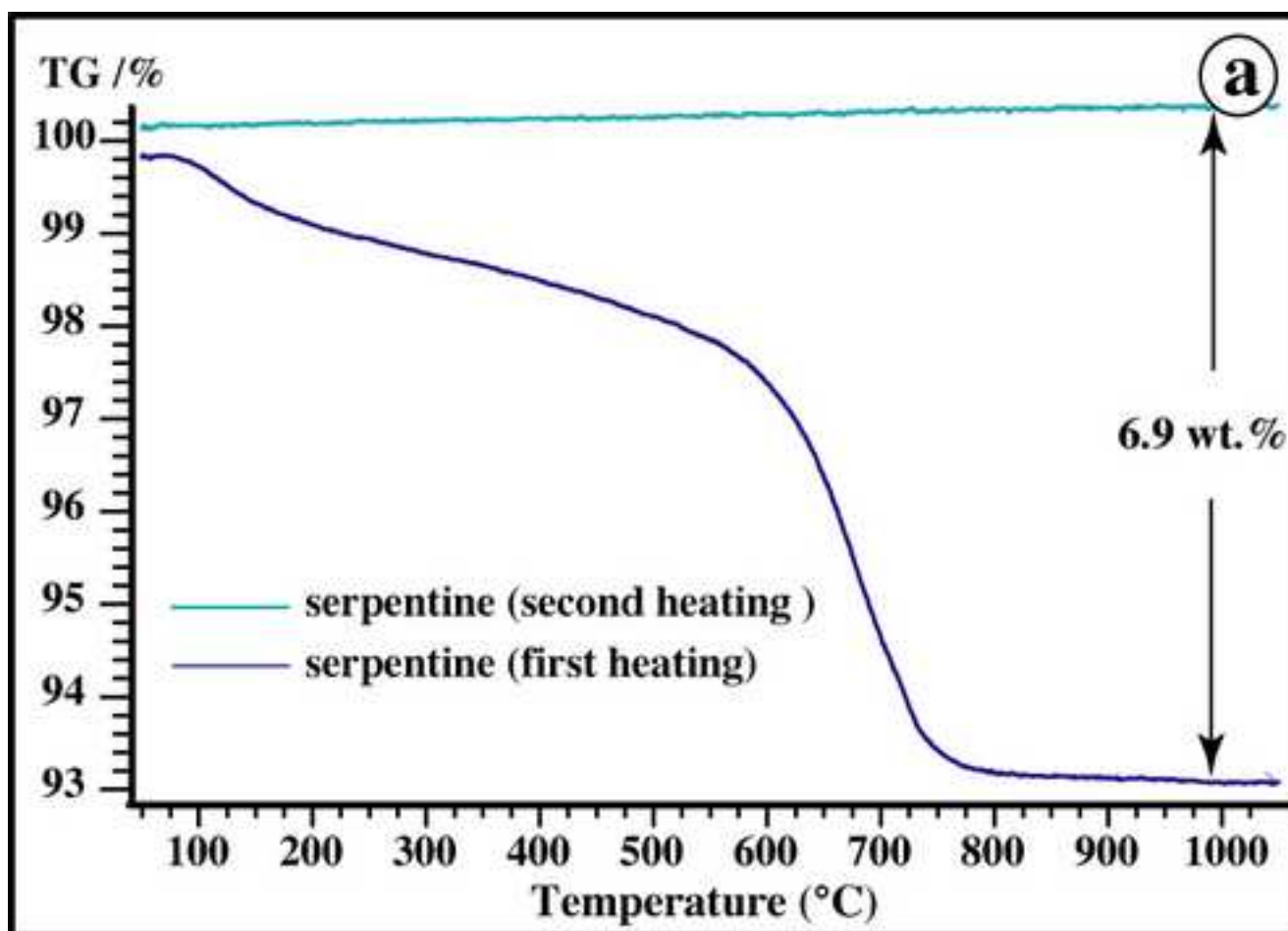


Figure12
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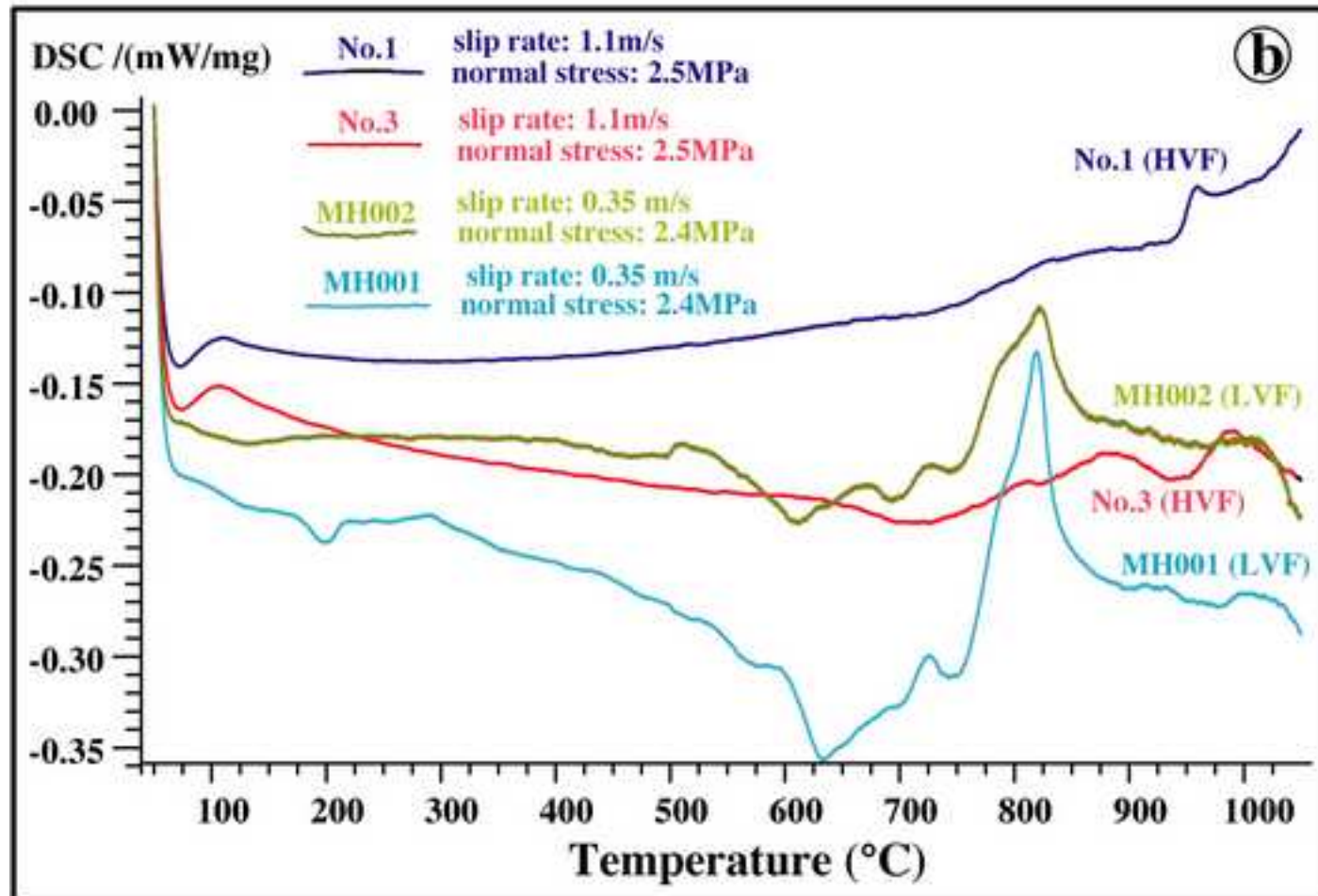
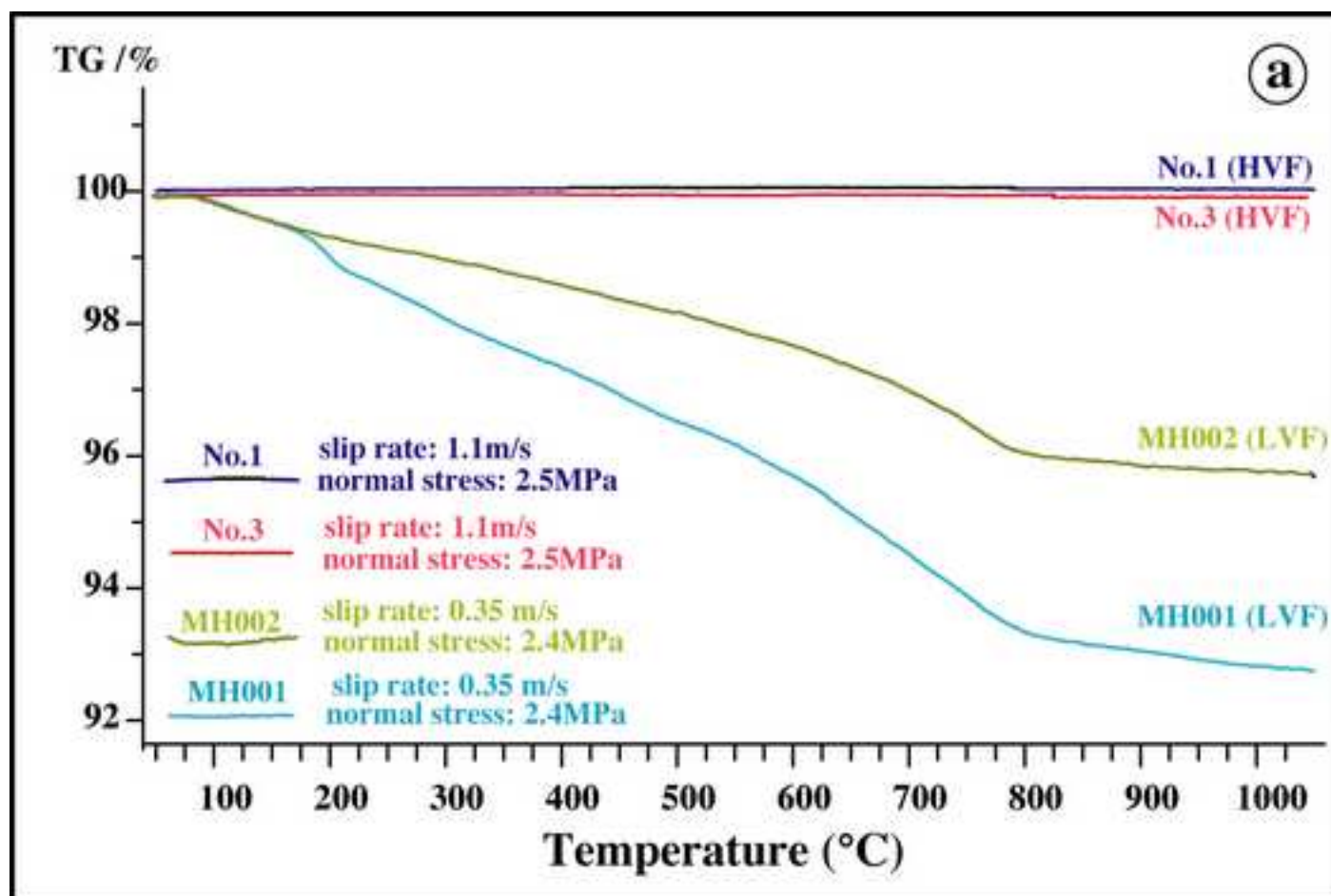


Table 1. Run products used for microscopic analyses.

<u>Run number</u>	<u>Normal stress (MPa)</u>	<u>Slip rate (m/s)</u>	<u>Analytical method</u>
No.1	2.5	1.10	TG-DSC analysis
No.3	2.5	1.10	TG-DSC analysis
MH001	2.4	0.35	TG-DSC analysis
MH002	2.4	0.35	TG-DSC analysis
MH003	2	1.10	FTIR analysis
MH004	9	1.10	FTIR analysis
MH012	5	1.31	FTIR analysis
HVF R001	6.5	1.10	Powder X-Ray analysis, EPMA
HVF R007	4	1.31	Powder X-Ray analysis, EPMA
THVR110	8.1	0.35	Frictional coefficient, shortening
HVF R012	5	1.31	Powder X-Ray analysis, EPMA

Table 2. Compositions of host rock (Host1, XRF) and serpentine (HS1-HS6, EPMA) and olivine (HO1-HO2, EPMA) materials in the host rock.

Sample No.	Host1	HS1	HS2	HS3	HS4	HS5	HS6	HO1	HO2
SiO ₂	41.580	41.322	40.648	41.944	40.068	41.422	42.828	40.576	41.185
TiO ₂	0.000	0.196	0.199	0.196	0.19	0.191	0.002	0.002	0.011
Al ₂ O ₃	1.098	3.076	3.104	2.919	3.610	3.224	1.777	0.175	0.800
FeO	9.904	3.063	2.887	3.674	3.268	3.225	7.074	6.327	9.335
MnO	0.228	0.137	0.170	0.181	0.14	0.166	0.127	0.140	0.181
MgO	45.908	36.694	38.194	37.997	37.141	38.112	43.460	51.206	46.294
CaO	1.242	0.106	0.114	0.106	0.095	0.088	1.260	0.127	0.526
Na ₂ O	0.034	0.086	0.108	0.093	0.125	0.111	0.074	0.010	0.036
K ₂ O	0.012	0.060	0.063	0.059	0.073	0.065	---	0.011	0.010
V	0.029	0.213	0.185	0.211	0.208	0.211	---	0.026	0.002
Cr ₂ O ₃	---	1.284	1.089	0.593	1.026	1.098	0.509	0.103	0.153
NiO	0.398	0.352	0.35	0.364	0.307	0.308	0.326	0.311	0.319
Total	99.973	86.588	87.11	88.337	86.249	88.219	97.438	99.013	98.852

XRF: X-ray fluorescence analysis; EPMA: Electron microprobe analysis;

FeO*: all Fe...; -not measured

Table 3. Chemical compositions of molten materials derived from the slip zone (M1-M8, EPMA).

Sample No.	M1	M2	M3	M4	M5	M6	M7	M8
SiO ₂	42.366	44.776	39.494	44.441	44.073	43.809	41.185	43.629
TiO ₂	0.015	0.012	0.001	0.025	0.022	0.003	0.011	0.000
Al ₂ O ₃	1.971	2.814	0.000	3.010	2.371	2.003	0.800	2.044
FeO*	7.144	7.860	11.243	7.465	7.289	6.704	9.335	7.391
MnO	0.081	0.077	0.207	0.095	0.141	0.080	0.181	0.103
MgO	43.176	40.053	48.68	38.99	42.037	44.681	46.294	42.209
CaO	1.595	2.020	0.014	2.275	1.953	1.138	0.526	1.422
Na ₂ O	0.126	0.126	0.003	0.159	0.139	0.097	0.036	0.076
K ₂ O	0.006	0.005	0.002	0.011	0.000	0.000	0.010	0.014
V	0.035	0.034	0.029	0.024	0.003	0.022	0.002	0.015
Cr ₂ O ₃	0.324	0.403	0.000	0.537	0.650	0.655	0.153	0.459
NiO	0.350	0.284	0.398	0.295	0.228	0.254	0.319	0.292
Total	97.188	98.464	99.558	97.328	98.908	99.446	98.852	97.654

EPMA: Electron microprobe analysis;

FeO*: all Fe

Table 4. Chemical compositions of serpentine minerals in the dehydration zones (Dh1-Dh9, EPMA).

Sample No.	Dh1	Dh2	Dh3	Dh4	Dh5	Dh6	Dh7	Dh8	Dh9
SiO ₂	47.462	45.76	46.881	45.543	44.184	46.035	44.832	44.243	46.742
TiO ₂	0.000	0.011	0.025	0.000	0.013	0.000	0.005	0.003	0.000
Al ₂ O ₃	3.198	3.107	4.020	2.549	3.123	2.768	3.813	3.116	1.998
FeO*	4.023	4.245	4.482	4.225	4.059	3.978	4.016	3.826	3.794
MnO	0.031	0.054	0.031	0.035	0.012	0.031	0.030	0.050	0.027
MgO	43.347	42.058	42.198	45.431	45.116	44.074	43.134	44.208	43.458
CaO	0.249	1.424	0.056	0.027	0.018	0.015	0.048	0.398	1.797
Na ₂ O	0.033	0.054	0.041	0.028	0.028	0.024	0.063	0.113	0.121
K ₂ O	0.000	0.010	0.004	0.000	0.012	0.002	0.001	0.005	0.013
V	0.024	0.020	0.034	0.018	0.025	0.018	0.040	0.030	0.001
Cr ₂ O ₃	0.419	0.236	0.870	0.463	0.889	0.476	1.441	1.605	0.621
NiO	0.151	0.160	0.134	0.181	0.214	0.165	0.223	0.157	0.163
Total	98.936	97.139	98.777	98.500	97.693	97.586	97.646	97.755	98.735

EPMA: Electron microprobe analysis;

FeO*: all Fe